



Modeling of Jupiter's stratosphere: new radiation code and impacts on the dynamics



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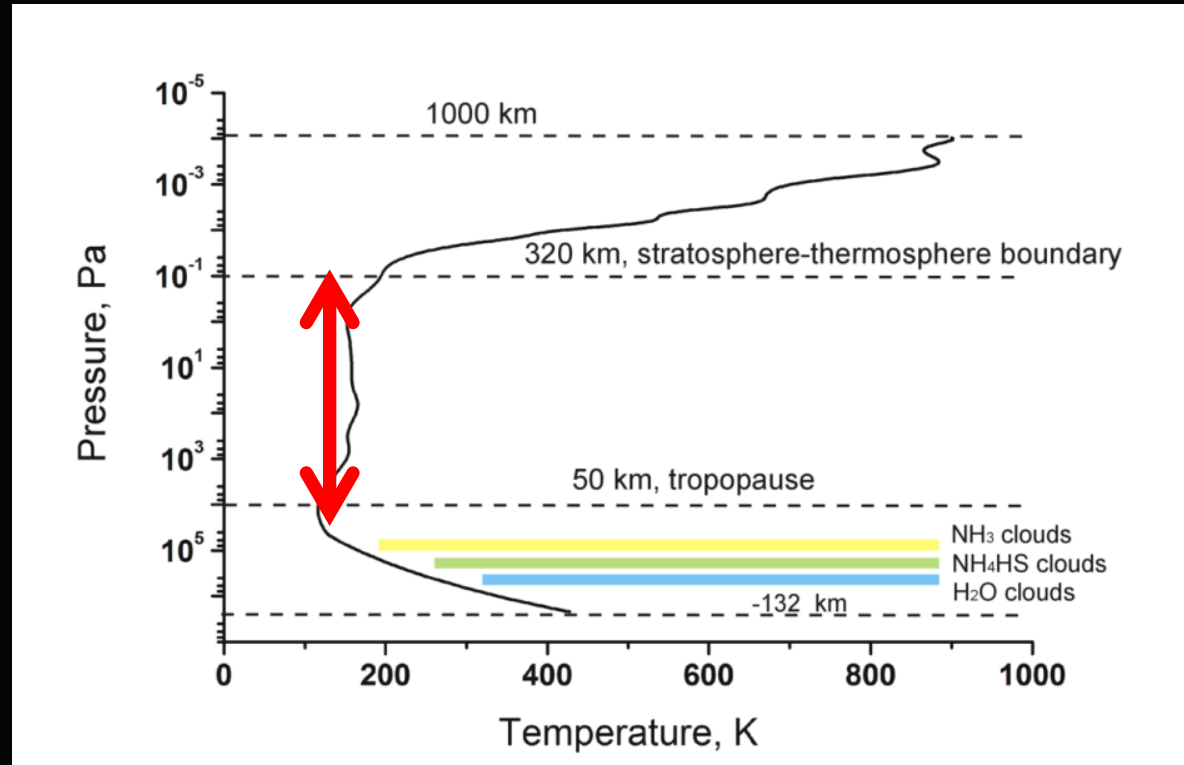
Max Planck Institute for Solar System Research



Atmosphere of Jupiter

Vertical structure: observed by Galileo Probe

- Thermosphere ($<10^{-3}$ hPa)
- Stratosphere ($10^2 \sim 10^{-3}$ hPa)
- Troposphere ($10^{4-5} \sim 10^2$ hPa)
 - With cloud layers
 - Driven by the internal heat source.



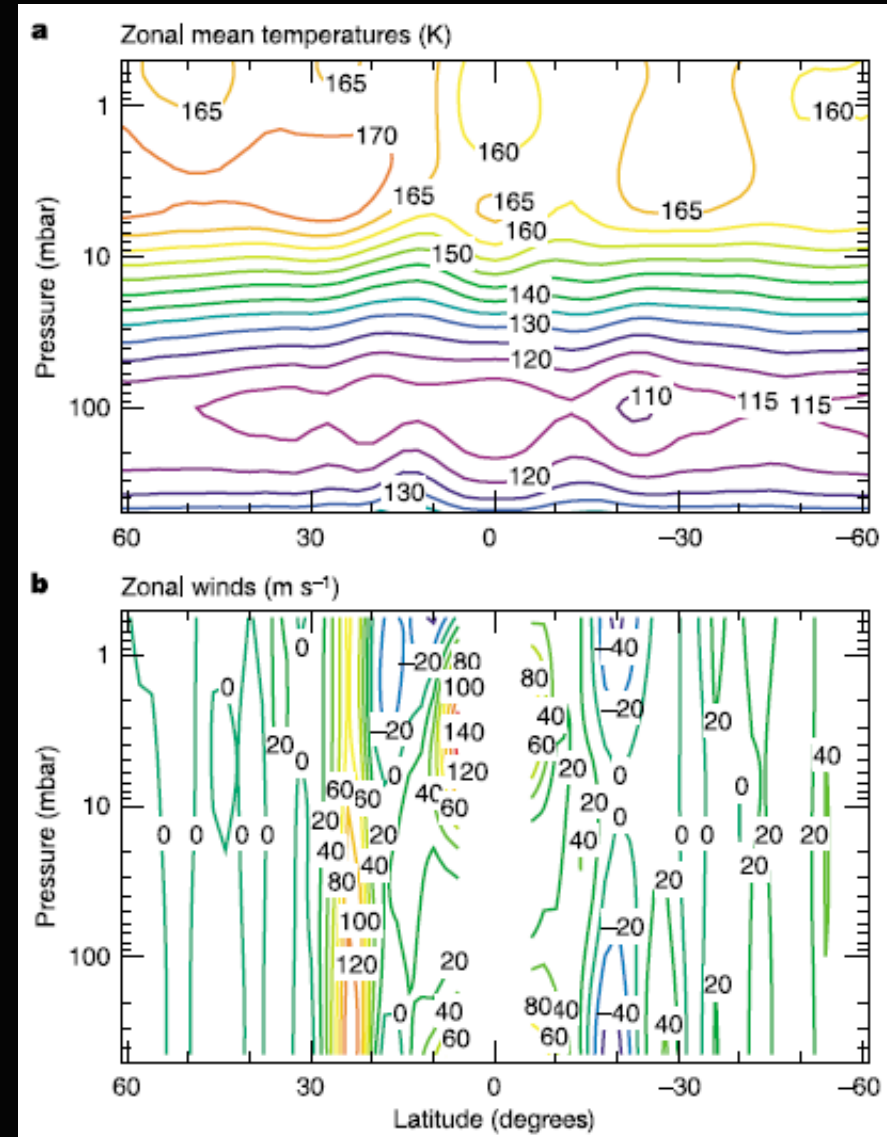
[Seiff et al., 1998]

Here we focus on the stratosphere.

Jupiter's stratosphere

- Affected by **radiative processes by molecules in stratosphere**, as well as **eddies** enhanced from the troposphere. (cf. troposphere: convection cell structures transport the energy and momentum)
- The estimation from the thermal wind equation and cloud tracking (for lower boundary wind speed) shows the existence of fast zonal wind jets of 60-140 m s⁻¹ at 23N and 5N.

Temperature and zonal wind fields observed by Cassini/CIRS



[Flasar et al., 2004]

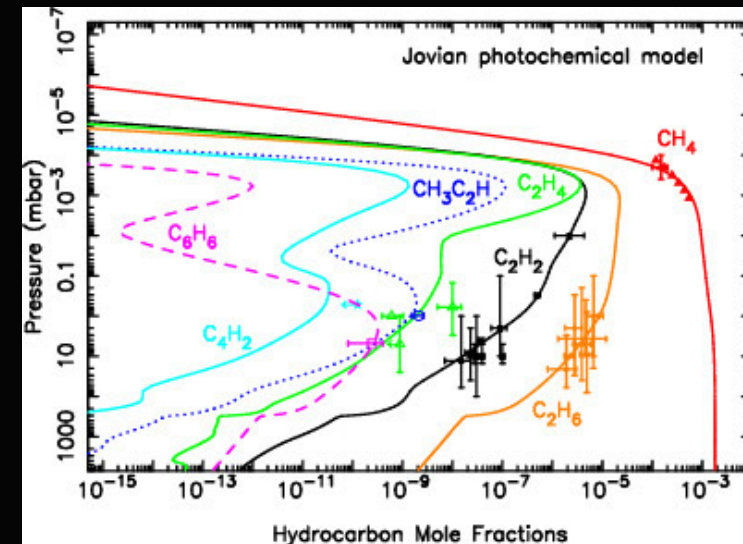
Radiative processes in Jupiter's stratosphere

- CH_4 : Absorber of the solar radiation
- CH_4 , C_2H_2 , C_2H_6 , collision-induced transitions of $\text{H}_2\text{-H}_2$ and $\text{H}_2\text{-He}$: Effective in the infrared cooling.

We have developed a band radiative transfer model for Jupiter's stratosphere for the fast and effective calculations in the GCM (correlated k -distribution approach).

- Here we show the vertical 1-D numerical results for
- heating/cooling rates and radiative balances
 - Radiative relaxation time
 - Radiative-convective equilibrium temperature

Mixing ratios of hydrocarbons from a photochemical model



[Moses et al., 2005]

Calculations

Coordinate of the band model

Band	IR(infrared) /SO(solar)	Wavenumber range [cm ⁻¹]	Molecules
1	IR	10-200	CH ₄ , H ₂ -H ₂ , H ₂ -He
2	IR	200-400	CH ₄ , H ₂ -H ₂ , H ₂ -He
3	IR	400-600	CH ₄ , H ₂ -H ₂ , H ₂ -He
4	IR	600-700	CH ₄ , C ₂ H ₂ , H ₂ -H ₂ , H ₂ -He
5	IR	700-860	C ₂ H ₂ , C ₂ H ₆ , H ₂ -H ₂ , H ₂ -He
6	IR	860-960	CH ₄ , C ₂ H ₆ , H ₂ -H ₂ , H ₂ -He
7	IR, SO	960-1200	CH ₄ , H ₂ -H ₂ , H ₂ -He
8	IR, SO	1200-1400	CH ₄ , H ₂ -H ₂ , H ₂ -He
9	IR, SO	1400-1700	CH ₄ , H ₂ -H ₂ , H ₂ -He
10	IR, SO	1700-2100	CH ₄ , H ₂ -H ₂ , H ₂ -He
11	SO	2100-3450	CH ₄ , H ₂ -H ₂
12	SO	3450-4800	CH ₄ , H ₂ -H ₂
13	SO	4800-6300	CH ₄ , H ₂ -H ₂
14	SO	6300-7800	CH ₄ , H ₂ -H ₂
15	SO	7800-9200	CH ₄ , H ₂ -H ₂
16	SO	9300-10800	CH ₄ , H ₂ -H ₂
17	SO	10800-11800	CH ₄ , H ₂ -H ₂

- Correlated k-distribution approach
- We made a table of k-distributions in **13 pressure grids** (log-equal interval between 10⁻³ and 10³ hPa), **3 temperature grids** (100, 150 and 200 K) for 17 wavenumber bands.

- The atmospheric composition of molecules (1000 ppmv of CH₄, 1 ppmv of C₂H₂, 10 ppmv of C₂H₆, 86.4 % of H₂, 13.6 % of He) is fixed in making the table.

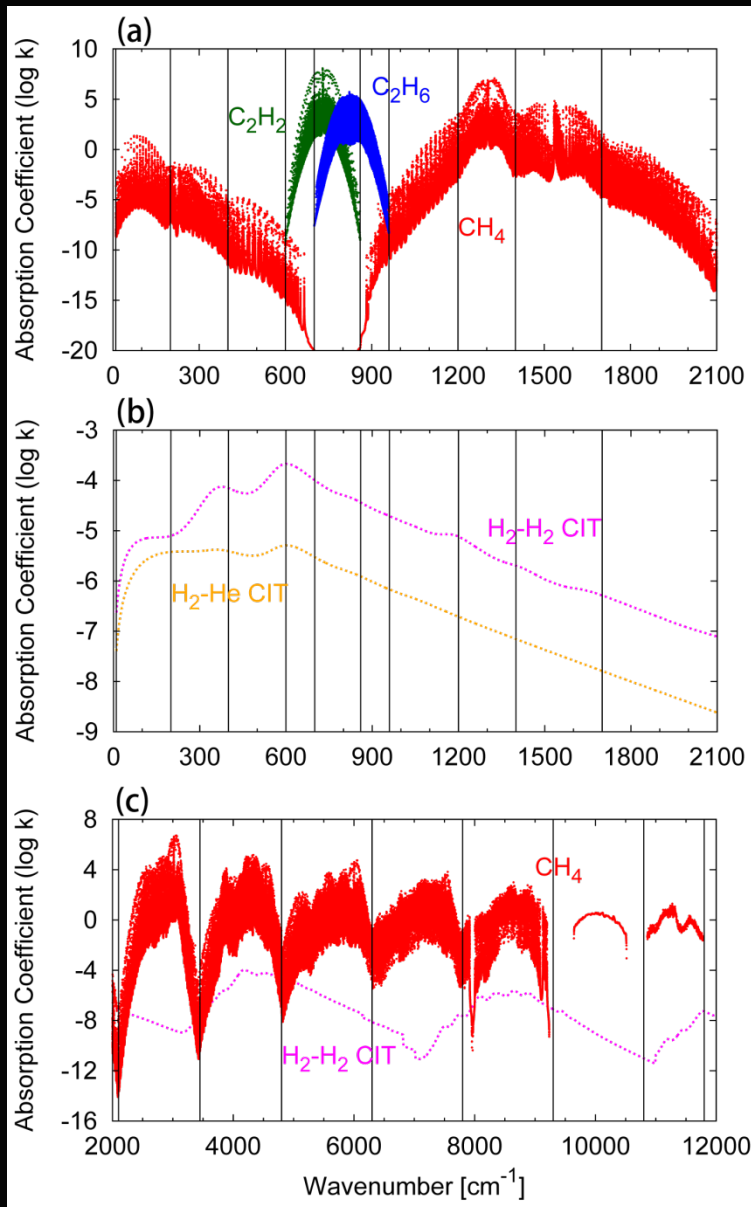
Calculations

Line spectra (1 hPa, 150K)

Molecules
(infrared)

Collision-
induced
transitions
(infrared)

For solar
absorption



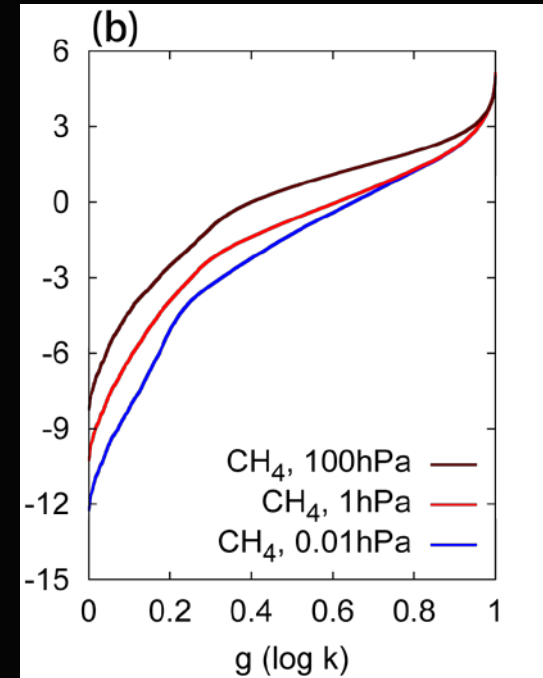
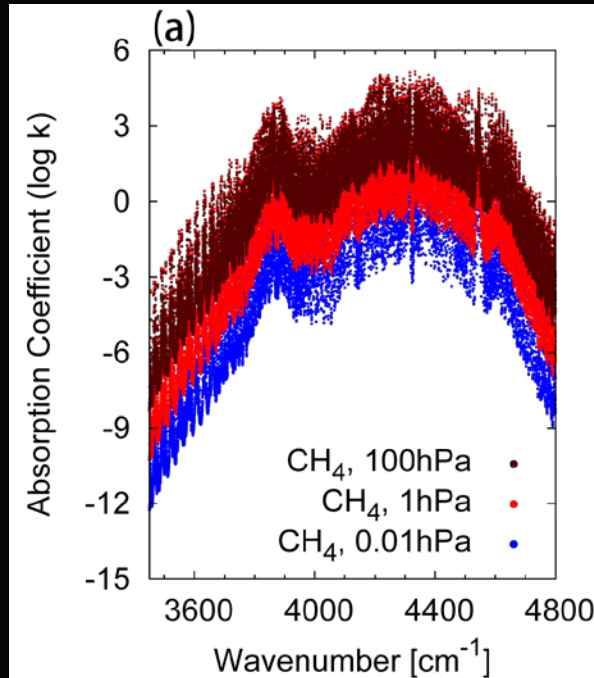
- Molecular lines of CH₄, C₂H₂ (600-860 cm⁻¹) and C₂H₆ (700-960 cm⁻¹): From HITRAN2008 [Rothman et al., 2009].
- Voigt profile is used for the calculation of line spectrum, with wing cutoff of 35 cm⁻¹ for all molecules.
- Collision-induced transitions of H₂-H₂ and H₂-He: From Borysow [2002] (H₂-H₂) and Borysow et al. [1988] (H₂-He).

Calculations

k-distribution [e.g. Liou, 2002]

CH₄ line spectra (3450-4800 cm⁻¹)

k-distribution of the line spectra



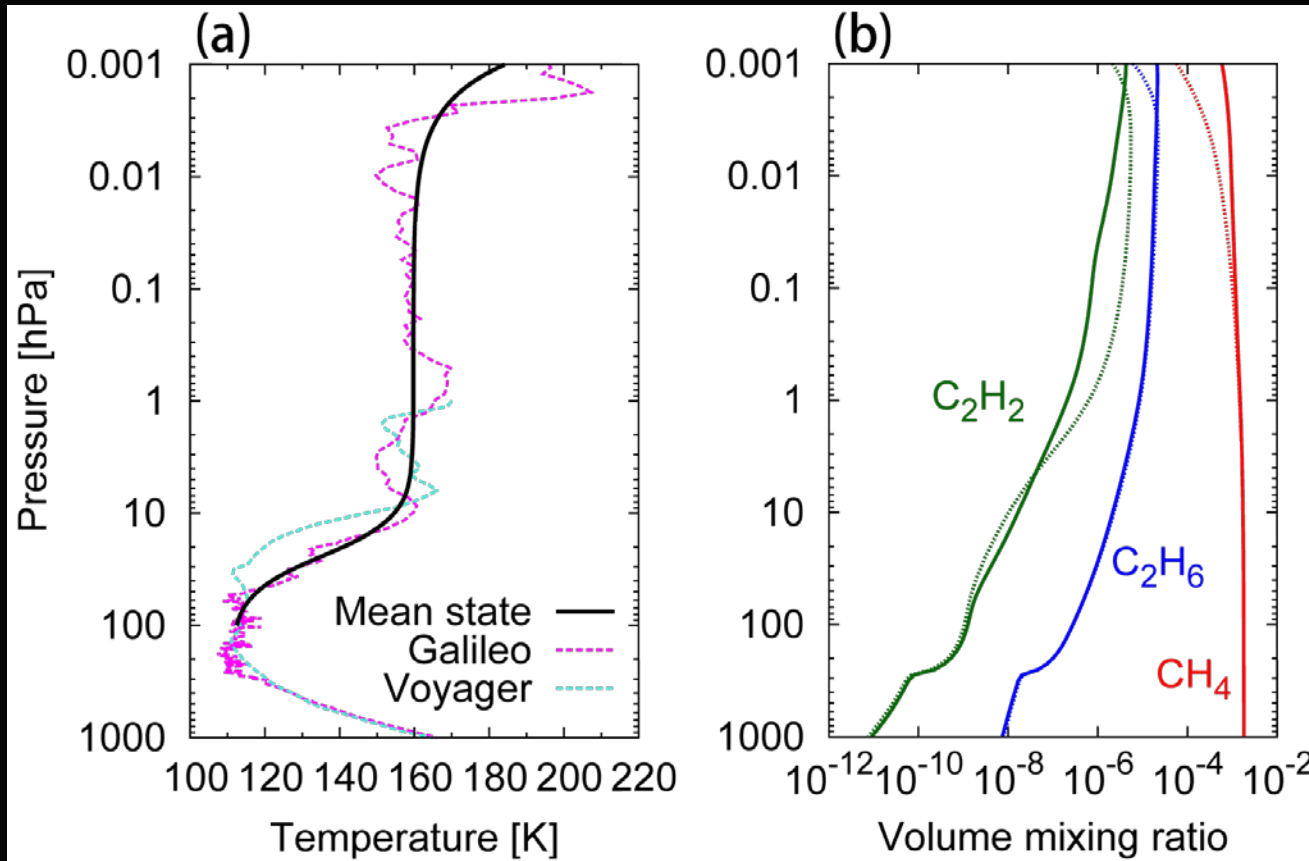
- For fast calculations of fluxes, the line spectrum in each band is ordered to be a monotone increasing function.
- The absorption and emission by molecules in each band are calculated with 12 k-distribution integration points per a molecule (144 points in the bands the lines of 2 molecules are overlapped).
- The effects of collision-induced transitions are added.

Calculations

Considered vertical profiles of temperature and composition

Temperature

Component

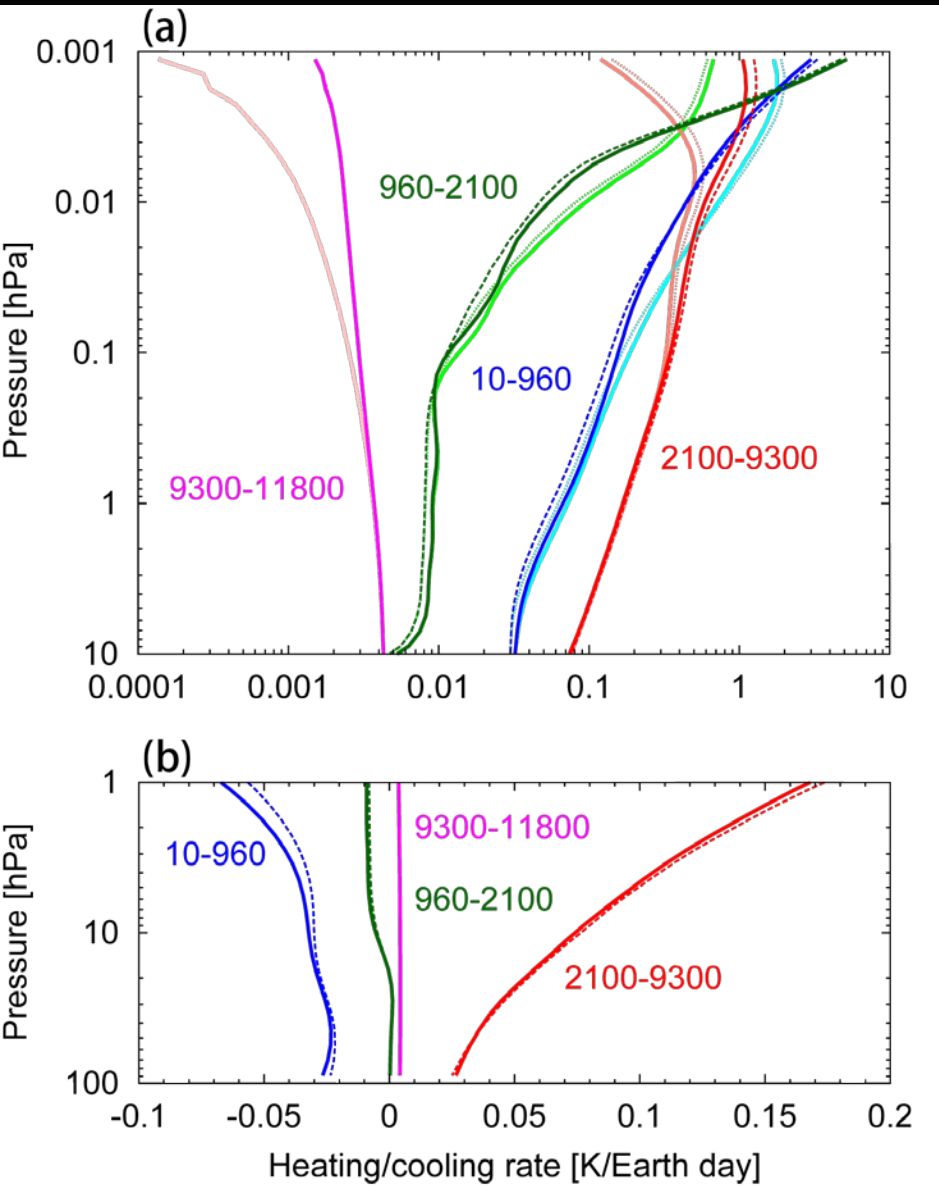


- Temperature: 'Mean state' from Galileo Probe observation [Yelle et al., 2001]
 - Component: From 1-D photochemical model [Moses et al., 2005]
- 2 kinds of results (Models A and C)

Results

Solid: Band
Dashed: Line-by-line

Heating/cooling rates



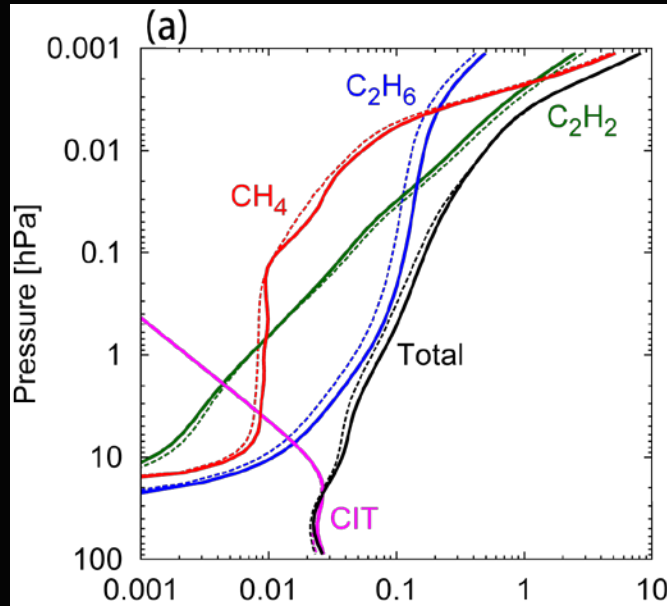
- Calculation of solar radiation: Assumed zenith angle of 0°
- Differences between band and line-by-line calculations are very small.
- Mid- and far-infrared radiation ($10\text{-}960\text{ cm}^{-1}$): Dominant for cooling below $\sim 2.5 \times 10^{-3}$ hPa.
- CH_4 infrared radiation ($960\text{-}2000\text{ cm}^{-1}$): Can be dominant for cooling above $\sim 2.5 \times 10^{-3}$ hPa, and very small effects below.
- Heating/cooling rates in upper stratosphere strongly depend on the composition.

Results

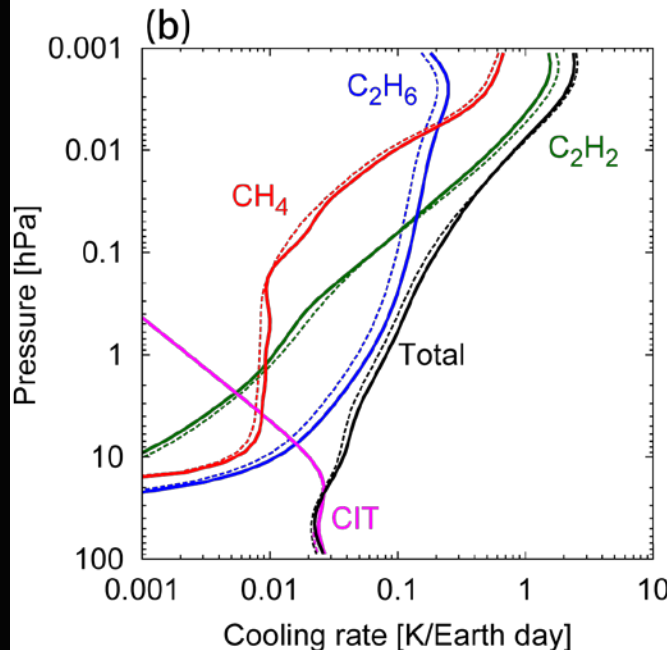
Solid: Band
Dashed: Line-by-line

Sensitivity of molecules (infrared cooling)

'Model A'
component



'Model C'
component



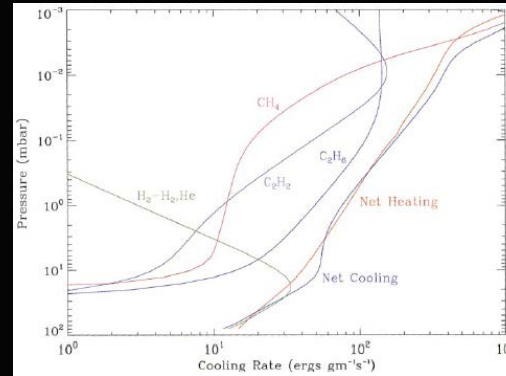
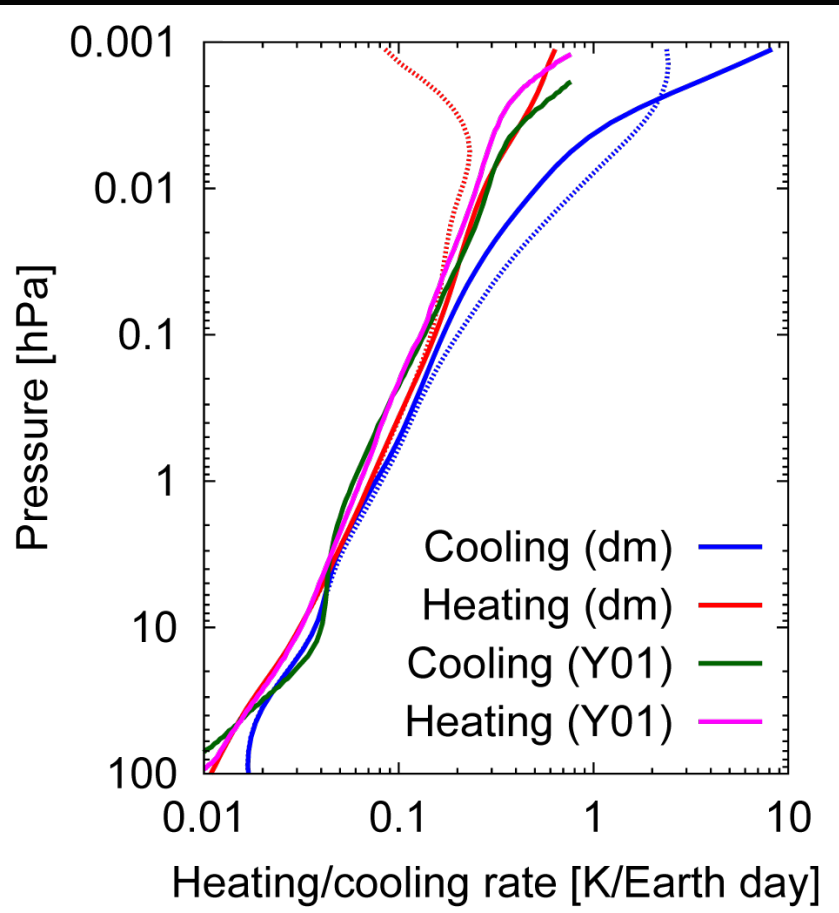
About the effect of cooling in
10-2100 cm⁻¹:

- C₂H₂ is dominant above ~0.03 hPa (up to ~3 K/day).
- C₂H₆ is dominant between 0.03-10 hPa (up to ~0.2 K/day in this height region).
- Collision-induced transitions are dominant below ~10 hPa (up to ~0.03 K/day).
- CH₄ can be dominant around the boundary to thermosphere, but its effect is small in most of the stratosphere.

Results

Total heating/cooling rate (in comparison with a preceding study)

Total day-mean heating&cooling rates
in comparison with Yelle et al. (2001)



[Yelle et al., 2001]

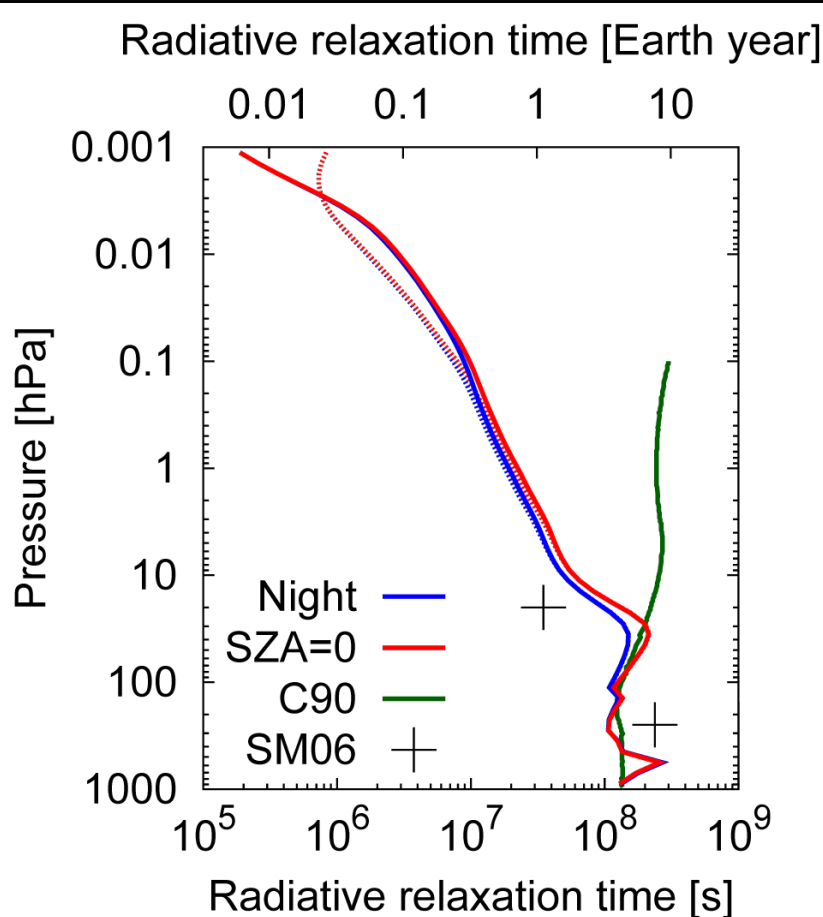
- Our calculations of day-mean net heating and cooling rates are in a good agreement with the results of Yelle et al. (2001), with radiative equilibrium.
- Above 0.1 hPa, our cooling rates exceed the heating rates, mainly due to stronger cooling by C₂H₂ in our model. (due to the lack of non-LTE effects...?)

Results

$$\tau = -\frac{\Delta T}{\Delta Q}$$

(From the difference of heating/ cooling rates for different temperature)

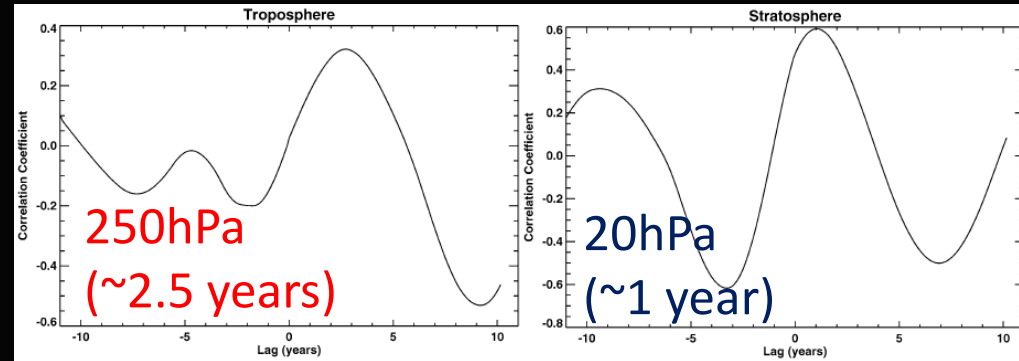
Daytime & nighttime,
Models A and C components



Radiative relaxation time

Cross correlation of subsolar latitude with the hemispheric temperature contrast (40°N-40°S) from IRTF observation (1979—2001)

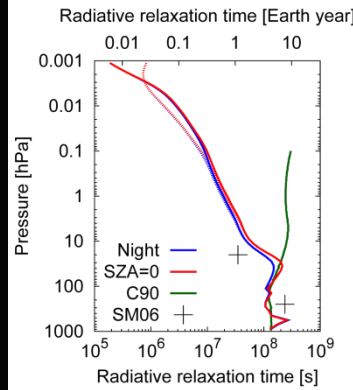
[Simon-Miller et al., 2006]



- The hemispheric temperature contrast lags the solar forcing longer in troposphere (~2.5 years) than in stratosphere (~1 year), which means the radiative relaxation time should be longer in troposphere.
- Our model shows qualitatively consistent results with the observation, while a preceding study [Conrath et al., 1990] does not.

Results

Equation of solar heating rate
[Conrath et al., 1990]

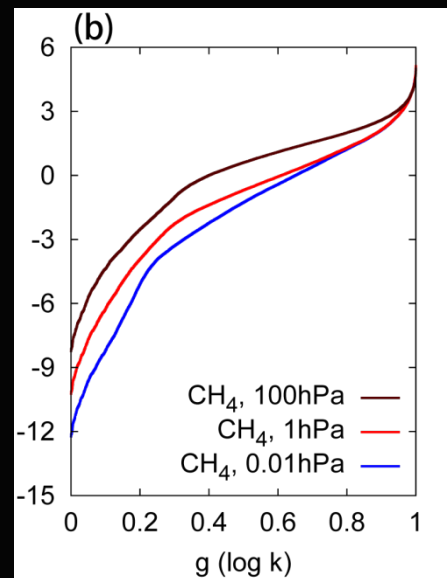
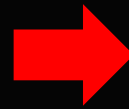
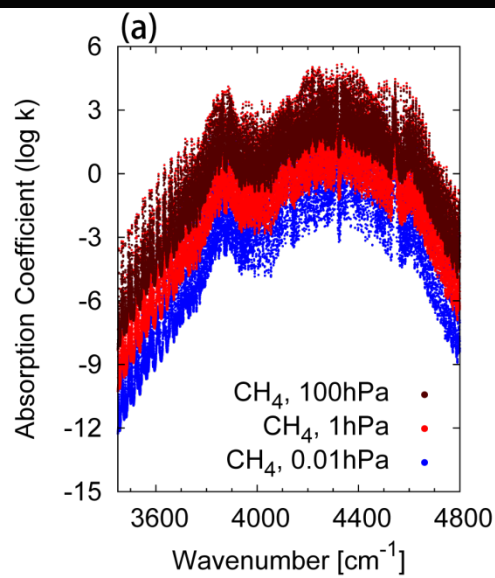


Radiative relaxation time

- The radiative relaxation time by Conrath et al. [1990] was shown to be longer in upper atmosphere, which contradicts the observations.
- It is because their model is simple and the heating/cooling rate is expressed to be proportional to the atmospheric density (pressure), which should underestimate the radiative effects in upper atmosphere.

$$Q_s = \rho g \mu_0 \sum_{i=1}^3 \frac{d \ln(\hat{p}_i N_1)}{dp} F_{\odot i} A_i \times \left[1 + \left(\frac{A_i d_i \mu_0}{2 S_i \gamma_i \hat{p}_i N_1} \right)^{1/2} \right]^{-1} + \rho g \frac{dN_1}{dp} \times (\bar{F}_{\odot a} \Delta v_a C_a e^{-C_a N_1 / \mu_0} + (\bar{F}_{\odot b} \Delta v_b C_b e^{-C_b N_1 / \mu_0})).$$

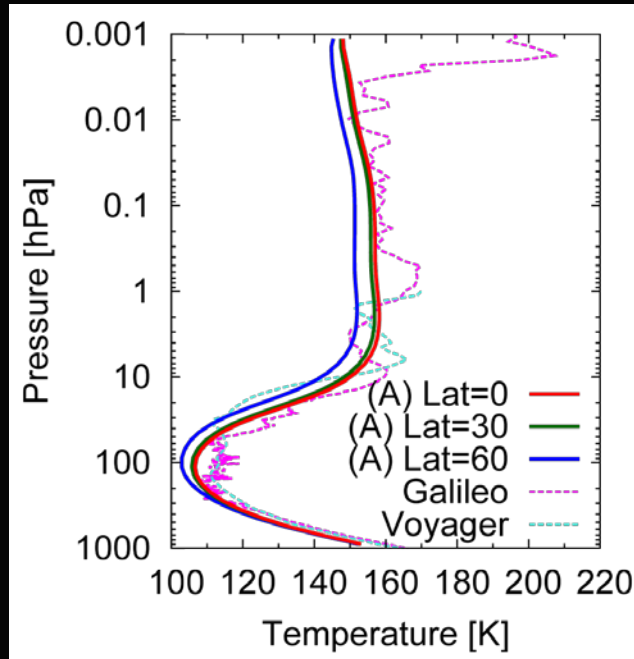
k-distribution



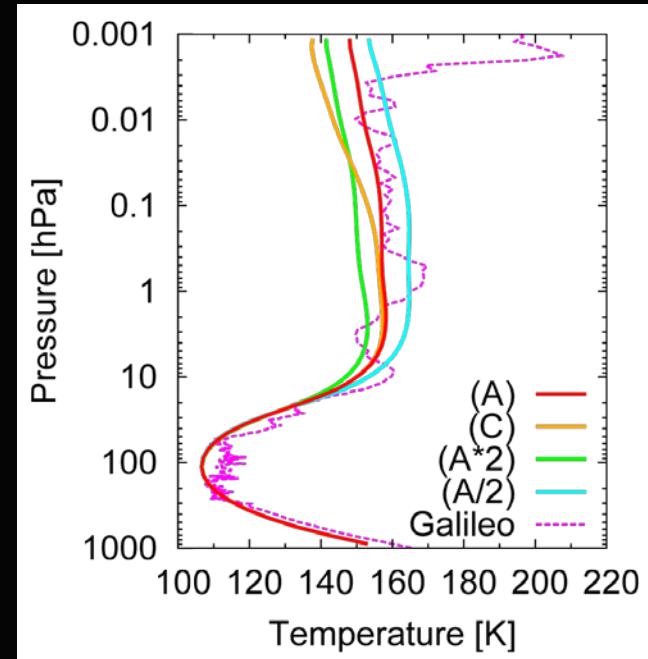
← At the peaks of spectra, the absorption coefficient becomes almost constant against the pressure. (except the peaks, proportional to pressure)

Results Radiative-convective equilibrium temperature

Different latitudes
(with observations)



Different components
(Models A and C, A with twice
more/less C_2H_2 and C_2H_6)



- Radiative-convective equilibrium temperature is close to the observed vertical profiles, except the upper troposphere (due to the lack of non-LTE effects...?)
- Small differences with different latitudes are indicated. (Note that the radiative effects of aerosols are not included)
- In upper stratosphere, it is sensitive to the components.

Development of the Jupiter stratospheric GCM

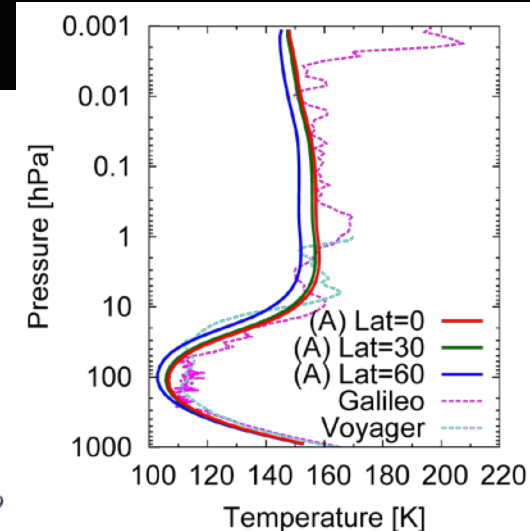
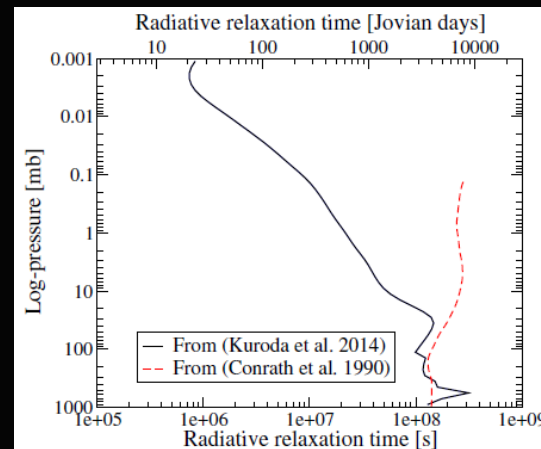
GCM descriptions

- Log-pressure coordinate in vertical
- 41 equally-spaced log-pressure levels in 0.01-1000 hPa (from cloud-top level to upper troposphere)
- Horizontal resolution of **240×180 grid points (1.5°×1°)** in longitude and latitude, correspondingly
- Radiative parameterization with Newtonian cooling, which relaxes the simulated temperature toward the prescribed equilibrium T_{eq}
- With different radiative relaxation time τ_{rad} : from Conrath et al. (1990) to this study

Newtonian cooling

$$F_T = (T_{eq} - T) / \tau_{rad}.$$

T_{eq} is defined from this result



Development of the Jupiter stratospheric GCM

Why the high-resolution is required?

Rossby radius of deformation

$$L_D = NH/f \propto T/gf,$$

Comparison with different planets [Showman et al., 2010; Sethunadh, 2014]

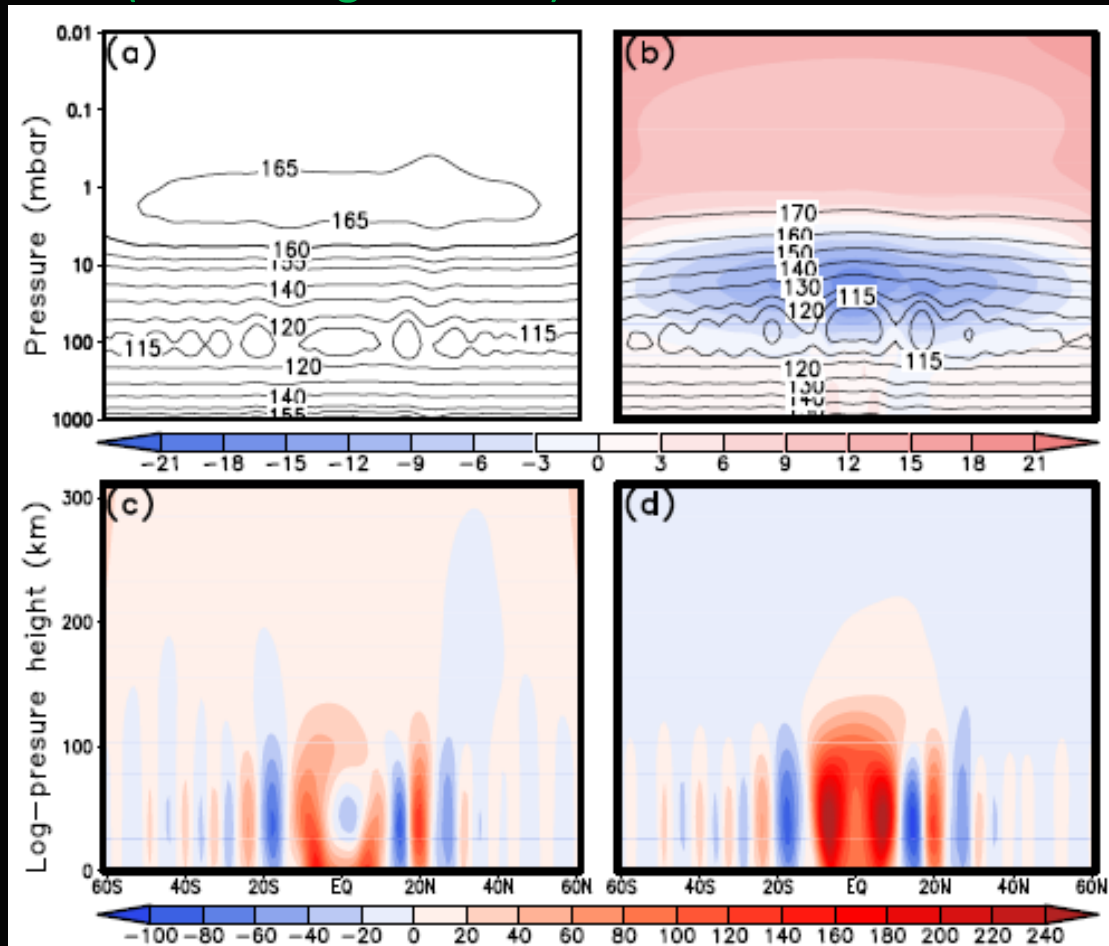
Planet	a (10 ³ km)	Ω (rad s ⁻¹)	gravity(ms ⁻²)	T_{eff} (K)	H(km)	U_c (ms ⁻¹)	L_D/a	L_β/a
Venus	6.05	3×10^{-7}	8.9	232	5	20	70	7
Earth	6.37	7.27×10^{-5}	9.82	255	7	20	0.3	0.5
Mars	3.396	7.1×10^{-5}	3.7	210	11	20	0.6	0.6
Titan	2.575	4.5×10^{-6}	1.4	85	18	20	10	3
Jupiter	71.492	1.7×10^{-4}	24.79	124	27	40	0.03	0.1
Saturn	60.268	1.65×10^{-4}	10.44	95	60	150	0.03	0.3
Uranus	25.56	9.7×10^{-5}	8.7	59	25	100	0.1	0.4
Neptune	24.76	1.09×10^{-4}	11.1	59	20	200	0.1	0.6

- The buoyancy force dominates the inertia for motions with the horizontal extent shorter than the Rossby radius of deformation.
- To simulate wave-mean flow interactions properly, GCMs must resolve motions shorter than the Rossby radius of deformation.
- Rossby radius of deformation is small for cold (small T), fast-rotating (large f), and massive (large g) planets like gas giants.

Development of the Jupiter stratospheric GCM

GCM results

$\tau_{\text{rad}} = 10^5$ s for all height (too strong in lower)
 τ_{rad} profile of this study



- It is seen that temperature adjusts closely to the prescribed T_{eq} under the strong radiative forcing.
- The zonal wind jets extend into the lower stratosphere and steeply decay with height.

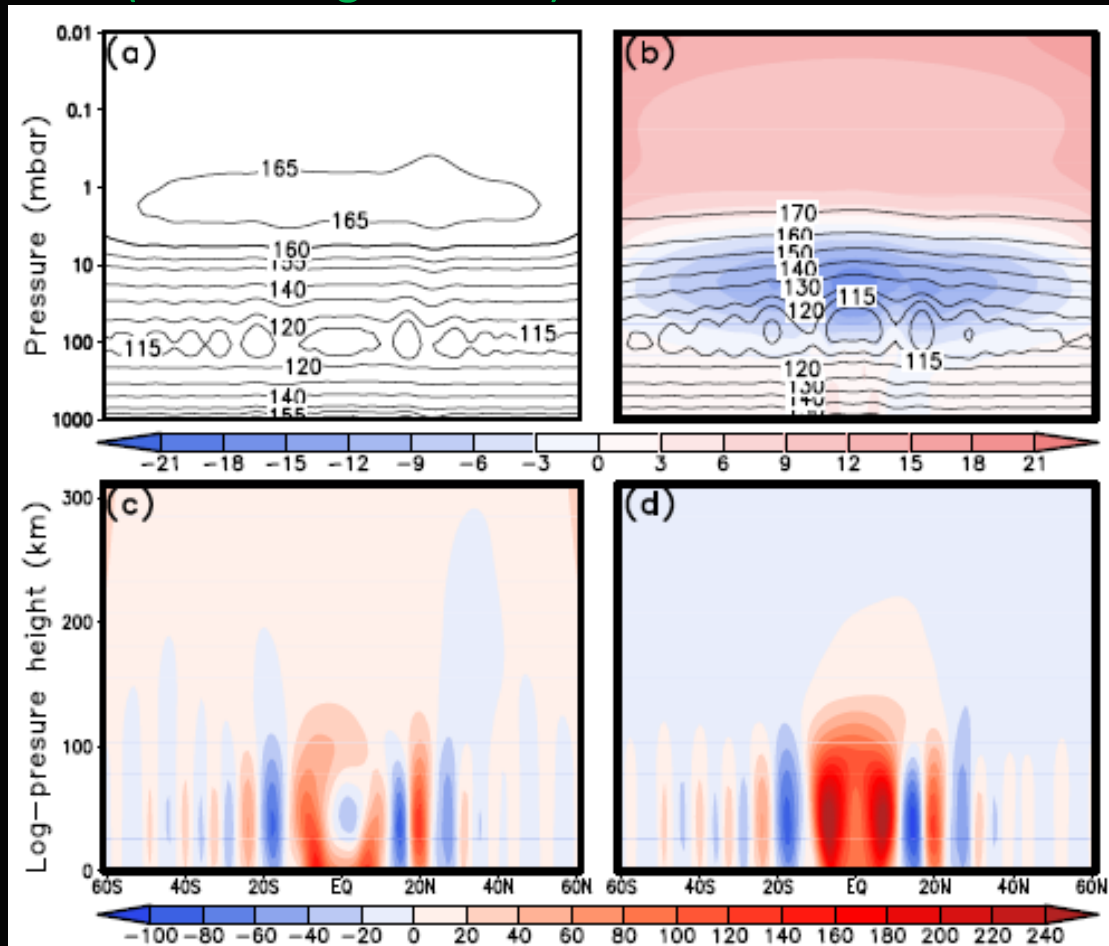
(Lower boundary wind velocity is defined from Cassini/VIMS cloud tracking)

Development of the Jupiter stratospheric GCM

GCM results

$\tau_{\text{rad}} = 10^5$ s for all height (too strong in lower)
 τ_{rad} profile of this study

- With larger τ_{rad} (corresponding to Conrath's) the calculation fails very rapidly with temperature dropping continuously.
- With $\tau_{\text{rad}} = 10^6$ s, simulations were very sensitive to the initial temperature disturbances.



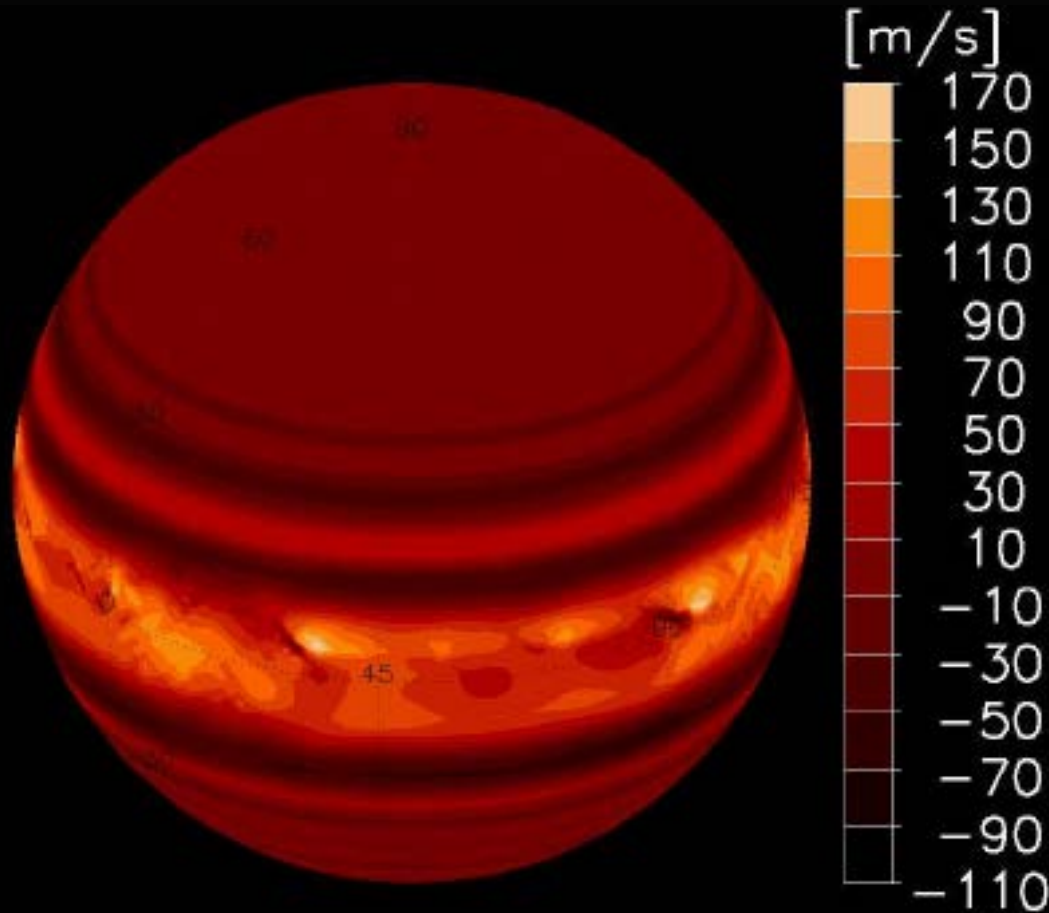
Temperature

Zonal wind

(Lower boundary wind velocity is defined from Cassini/VIMS cloud tracking)

Development of the Jupiter stratospheric GCM

GCM results



Zonal wind distribution at 30hPa using the vertical profile of radiative relaxation time in this study

Now the implementation of the radiation code of this study into the GCM is ongoing!

(Lower boundary wind velocity is defined from Cassini/VIMS cloud tracking)

Summary

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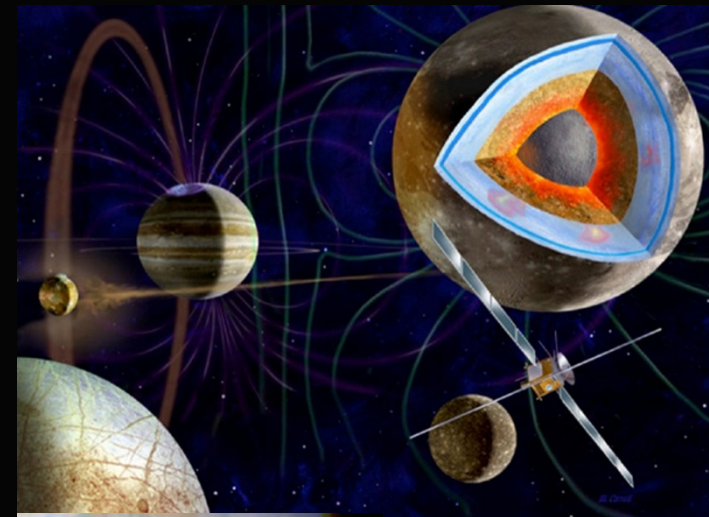
“Parameterization of radiative heating and cooling rates in the stratosphere of Jupiter”.

- A fast and effective band model for Jupiter’s stratosphere was developed, calculating the heating/cooling rates in a good accuracy in comparison with the line-by-line calculations.
- The band model showed radiative equilibrium in the middle of Jupiter’s stratosphere. In the upper stratosphere, the heat balance is very sensitive to the mixing ratios of hydrocarbons.
- It also showed that the radiative relaxation time becomes shorter in upper atmosphere, which is consistent with the observations [Simon-Miller et al., 2006] and corrects the theoretical error in the preceding study [Conrath et al., 1990].
- Radiative-convective equilibrium temperature was calculated for different latitudes and composition. In low-latitude region, it is close to the observed temperature profiles.
- Now we are starting the study with a Jupiter’s stratospheric GCM which requires a high resolution. Implementation of this radiation code to the GCM is now ongoing.

This study is in connection with

JUICE-SWI (Sub-Millimetre Instrument)

- The main objective of a sub-millimetre wave instrument is to investigate **the structure, composition and dynamics of the middle atmosphere of Jupiter** and exospheres of its moons, as well as thermophysical properties of the satellites surfaces. (from Yellow Book)
- JUICE-SWI is highly sensitive to CH_4 , H_2O , HCN , CO and CS in Jupiter's stratosphere.
- From CH_4 molecular lines, vertical temperature profiles and wind velocities can be measured.
- CO and CS , which are chemically stable, can be used as tracers for the investigations of atmospheric flows (general circulation and dynamical processes).



Collision of Shoemaker-Levy 9
[HST, 1994]:
Origin of H_2O , CS , CO and HCN ?

Details will be given by Dr. P. Hartogh (tomorrow 17:20-)!