

# Stability of Atmospheric Redox States of Early Mars Inferred from Time Response of the Regulation of H and O Losses

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## Abstract:

Atmospheric losses including escapes to space and depositions to the surface play an essential role in the evolution of the Martian surface environment. Especially, a ratio of total losses of hydrogen and oxygen from the atmosphere is crucial to determine its atmospheric redox state. In the condition that H and O losses originate from H<sub>2</sub>O, an atmospheric redox state remains the same if the ratio is 2:1, otherwise it is driven into oxidizing or reducing. It was shown by McElroy. (1972) that Jeans escape flux of H and H<sub>2</sub> and nonthermal escape flux of O were regulated to be in the ratio of 2:1 in a converged model of present-day Mars, which is called “self-regulation”. Whether or not the self-regulation works in real atmospheres depends on its timescales, but time responses of the self-regulation are not well understood in different atmospheric conditions.

Here we study time responses of the self-regulation in different atmospheric conditions and discuss the stability of atmospheric redox states. We use a 1D time-dependent photochemical model for various atmospheric conditions and parameters, such as atmospheric CO<sub>2</sub> pressure, surface temperature and O escape rate.

We find that the self-regulation timescale is essentially controlled by the net redox balance ( $pO_x [\text{mbar}] = 2pO_2 - pCO - pH_2$ ) in a converged state. The timescale gets longer as  $|pO_x|$  increases, which suggests that redox-neutral atmospheres have the shortest timescale. We also find that the self-regulation can be classified into two regimes. First regime is the same as the one explained by Liu and Donahue. (1976), which tends to work in oxidizing atmospheres ( $pO_x > 0$ ) including present-day Mars in a way that H escape changes to reach the regulated state following a change in H<sub>2</sub> transportation from the lower to upper atmosphere. Second one is likely to work in thicker and reducing atmospheres ( $pO_x < 0$ ) over a relatively long timescale. The regulation occurs dominantly by changes in CO abundance in the lower atmosphere. These results imply that thicker atmospheres in early Mars are less redox-stable than present-day Mars. Our model calculations also indicate that CO-dominated atmosphere of about 100 mbar might be possible at ~3 Ga. We finally discuss the redox stability of H<sub>2</sub>-rich CO<sub>2</sub> atmosphere of early Mars.

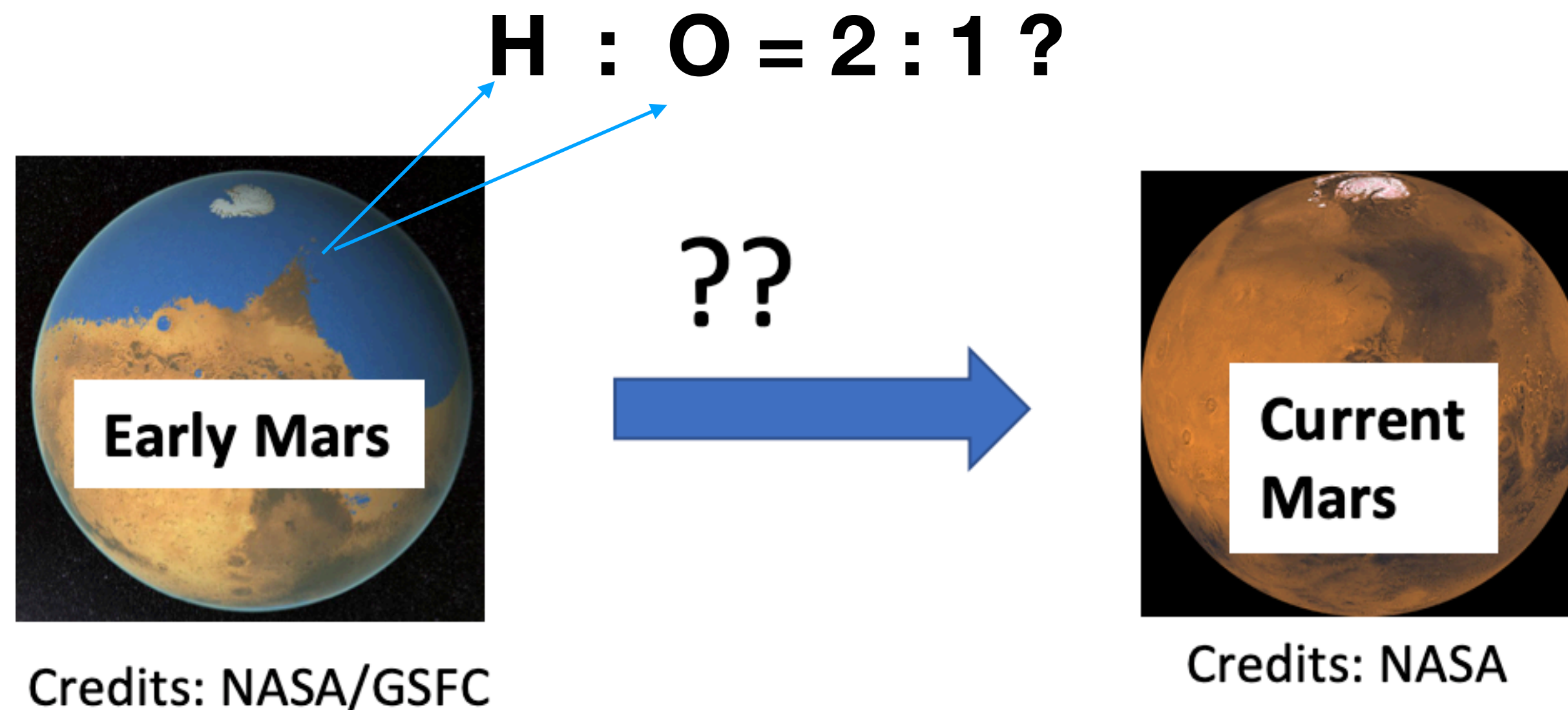
# **Stability of atmospheric redox states of early Mars inferred from time response of the regulation of H and O losses**

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# The effect of the ratio of H & O losses

- How did H<sub>2</sub>O on early Mars go away?



- The ratio of total H and O outfluxes

$\Phi_{total\ H} > 2\Phi_{total\ O}$   $\longrightarrow$  Oxidizing atmosphere

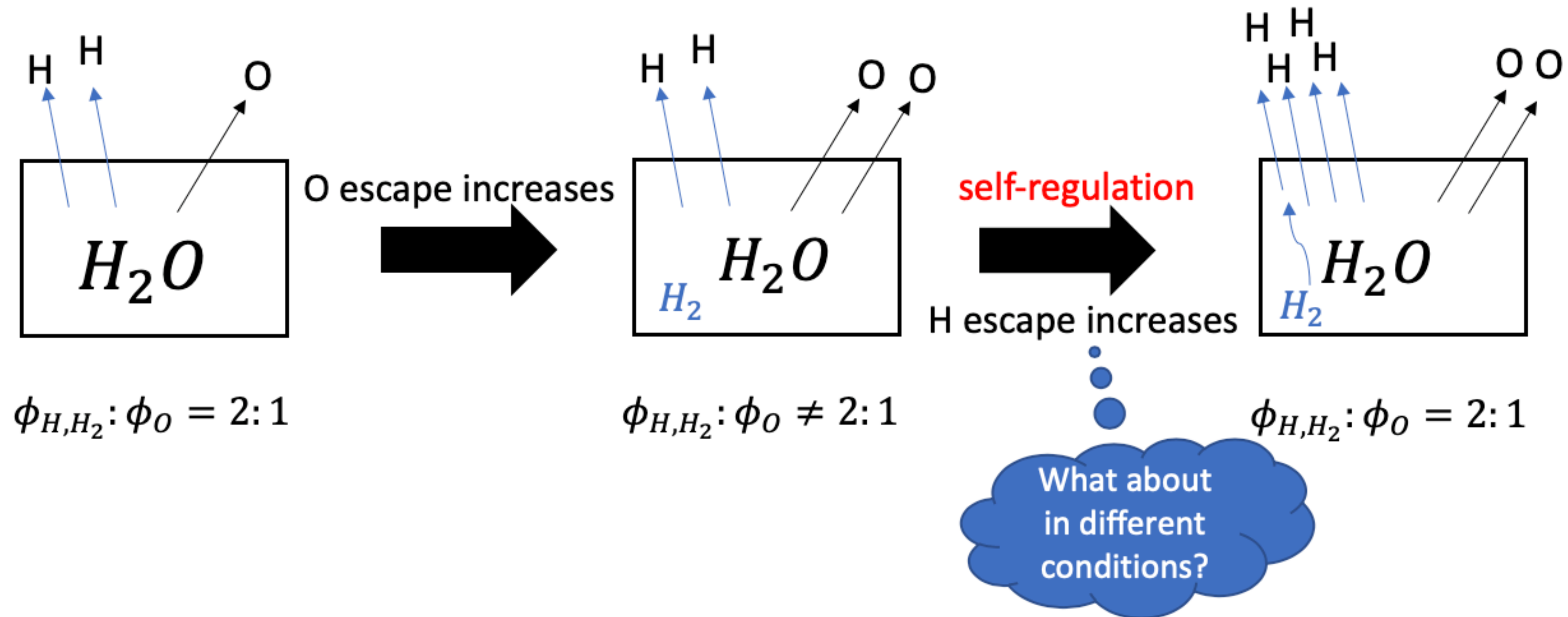
$\Phi_{total\ H} < 2\Phi_{total\ O}$   $\longrightarrow$  Reducing atmosphere

- Remain the same?
- More reducing?
- More oxidizing?

# What is Self-regulation of H & O losses?

- Jeans escape flux of H and H<sub>2</sub> and nonthermal escape flux of O are regulated to be in the ratio of 2:1 in a converged model of present-day Mars.

[McElroy, 1972; Liu & Donahue, 1976]

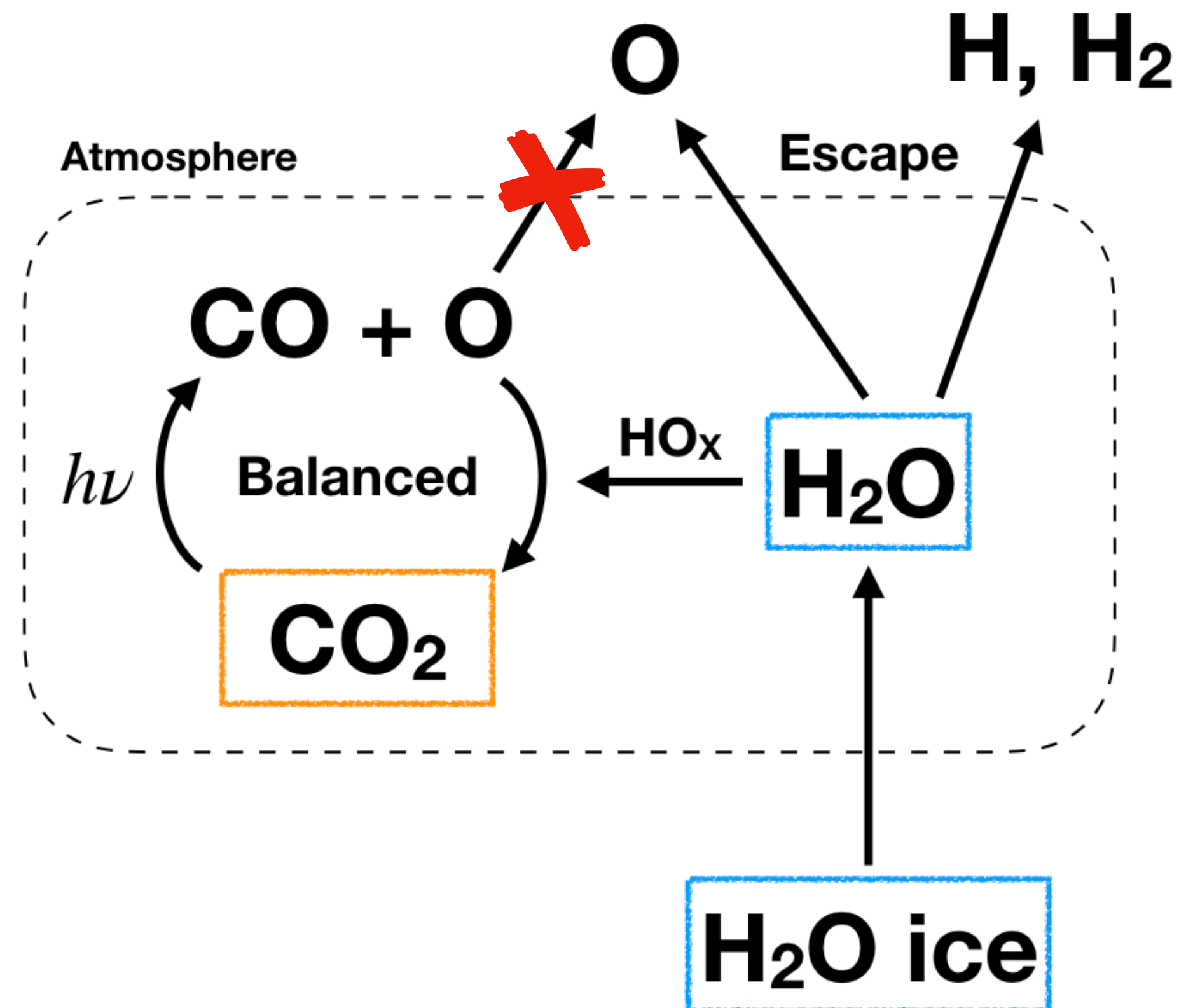


**Self-regulation is not well understood in different conditions.**

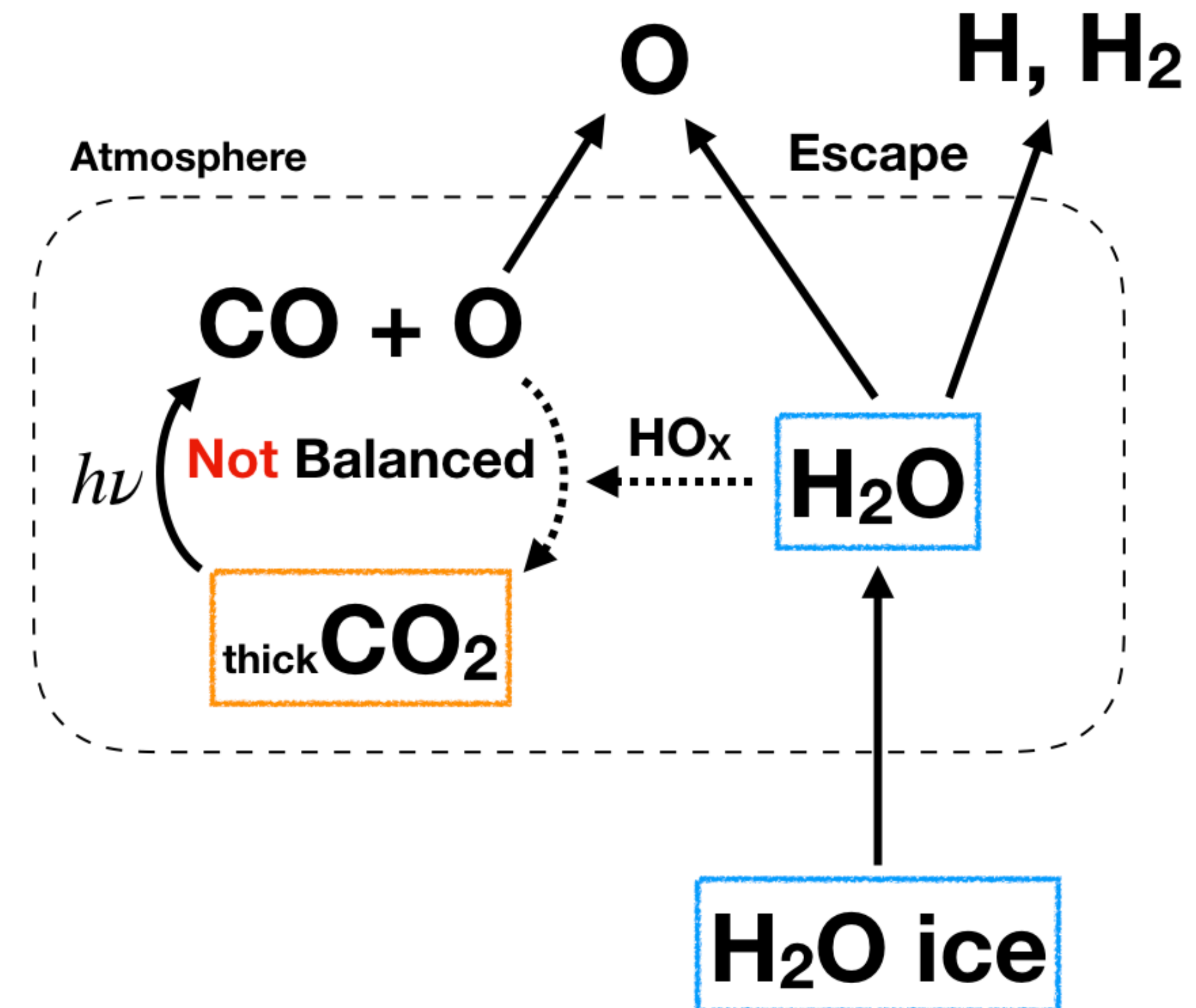
# Two essential factors in other atmospheric conditions

- ① **CO<sub>2</sub> stable or CO-runaway ?**
- ② Self-regulation timescale

CO<sub>2</sub> stable



CO-runaway





# Two essential factors in other atmospheric conditions

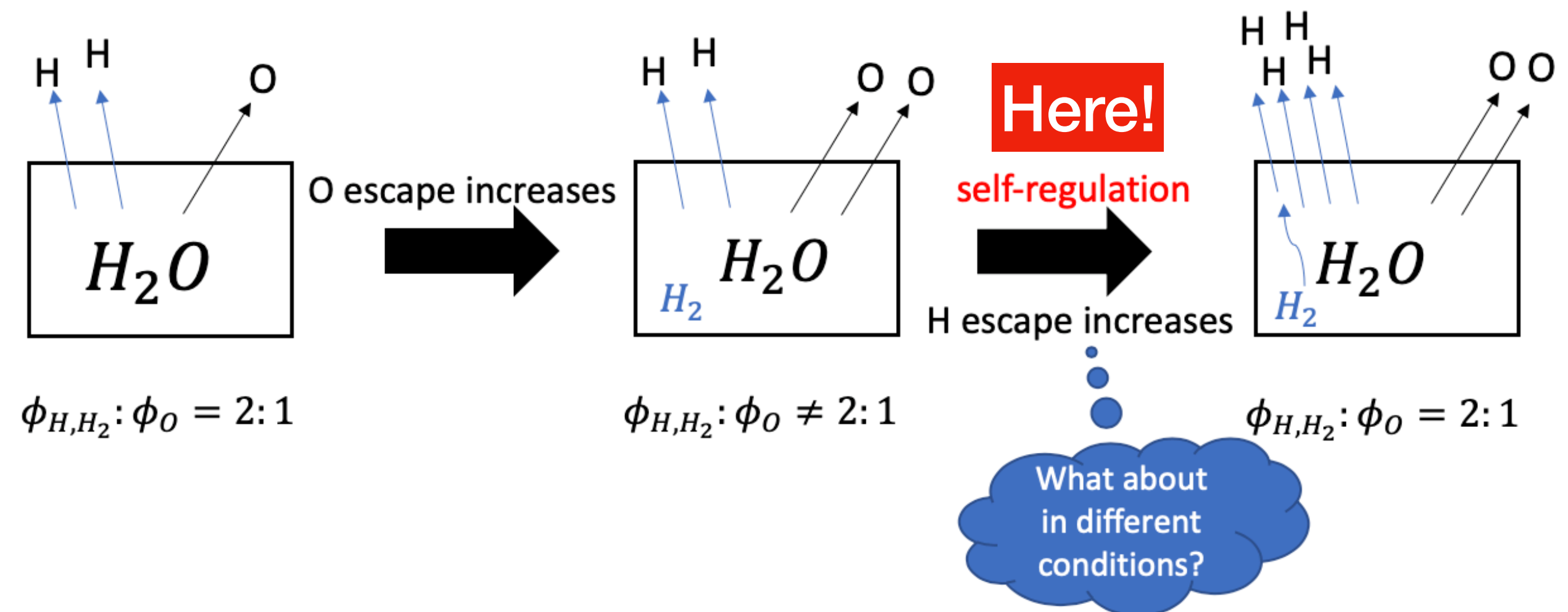
① CO<sub>2</sub> stable or CO-runaway ?

② **Self-regulation timescale**

Shorter timescale

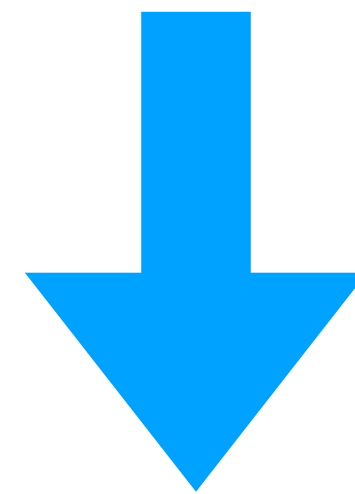


Redox state: more stable



# Purpose

To investigate time responses of the self-regulation in different atmospheric conditions from present-day Mars



**Timescales of the self-regulation**

**Stability of atmospheric redox states in early Mars**

# 1D-Photochemical model

modified after Chaffin et al. (2017)

## ❖ Basic model information

- Initial  $p\text{CO}_2 = 6.3\text{mbar} \sim 3\text{bar}$
- 13 species(C,H,O-bearing species)
- 52 reactions

## ❖ Boundary condition

Upper	O	$\Phi_{\text{O}_{\text{esc}}} = 1.2 \times 10^8 \text{ cm}^{-2} \text{ s}^{-1}$
	H, H <sub>2</sub>	Jeans' Escape
Lower	reactive species	deposition velocity $v_{\text{dep}} = 0.02 \text{ cm s}^{-1}$ [Zahnle et al., 2008]
	CO <sub>2</sub>	Not fixed

Reactive species: O<sub>3</sub>, H<sub>2</sub>O<sub>2</sub>, O, H, OH, HO<sub>2</sub>, O<sub>1</sub>D

## ❖ Temperature profile

$T_{\text{surf}} = 200\text{-}280\text{K}$

Ex) CO<sub>2</sub>:10mbar

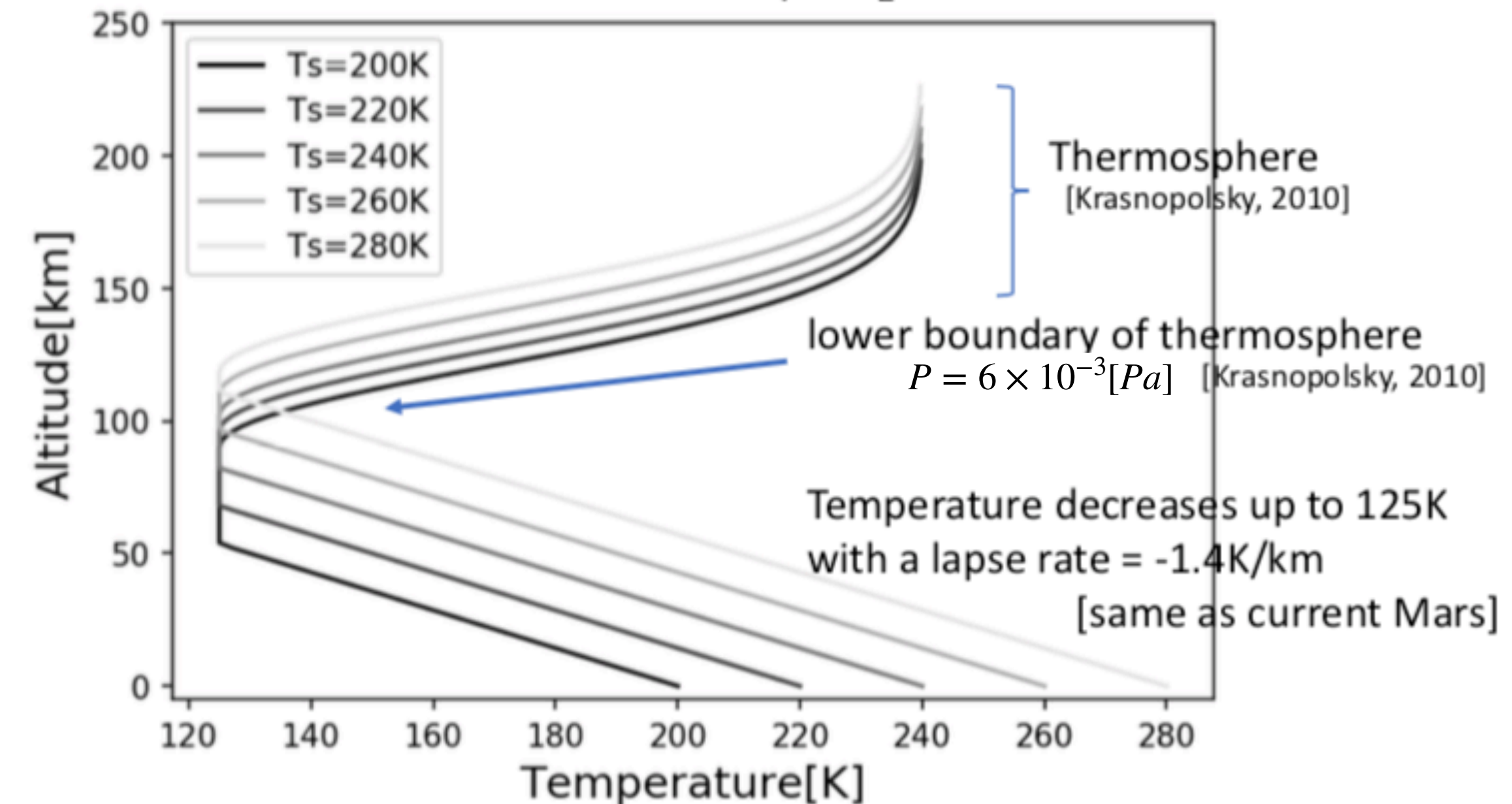


Fig. Vertical temperature profile.

## ❖ Water vapor profile

- Relative humidity is fixed at 0.22 near the surface
- Following the saturation water vapor curve above



# Definition of Self-regulation timescale

At time zero, oxygen escape increases by a factor of two from a steady state, we define "timescale" of self-regulation as how long it takes to reach  $\varepsilon < 0.01$ .

At time=0

$$\Phi_{O_{esc}} \rightarrow 2\Phi_{O_{esc}}$$

②

Self-regulated state

$$\varepsilon = \left| 2 - \frac{\Phi_{\text{total H}}}{\Phi_{\text{total O}}} \right| < 0.01$$

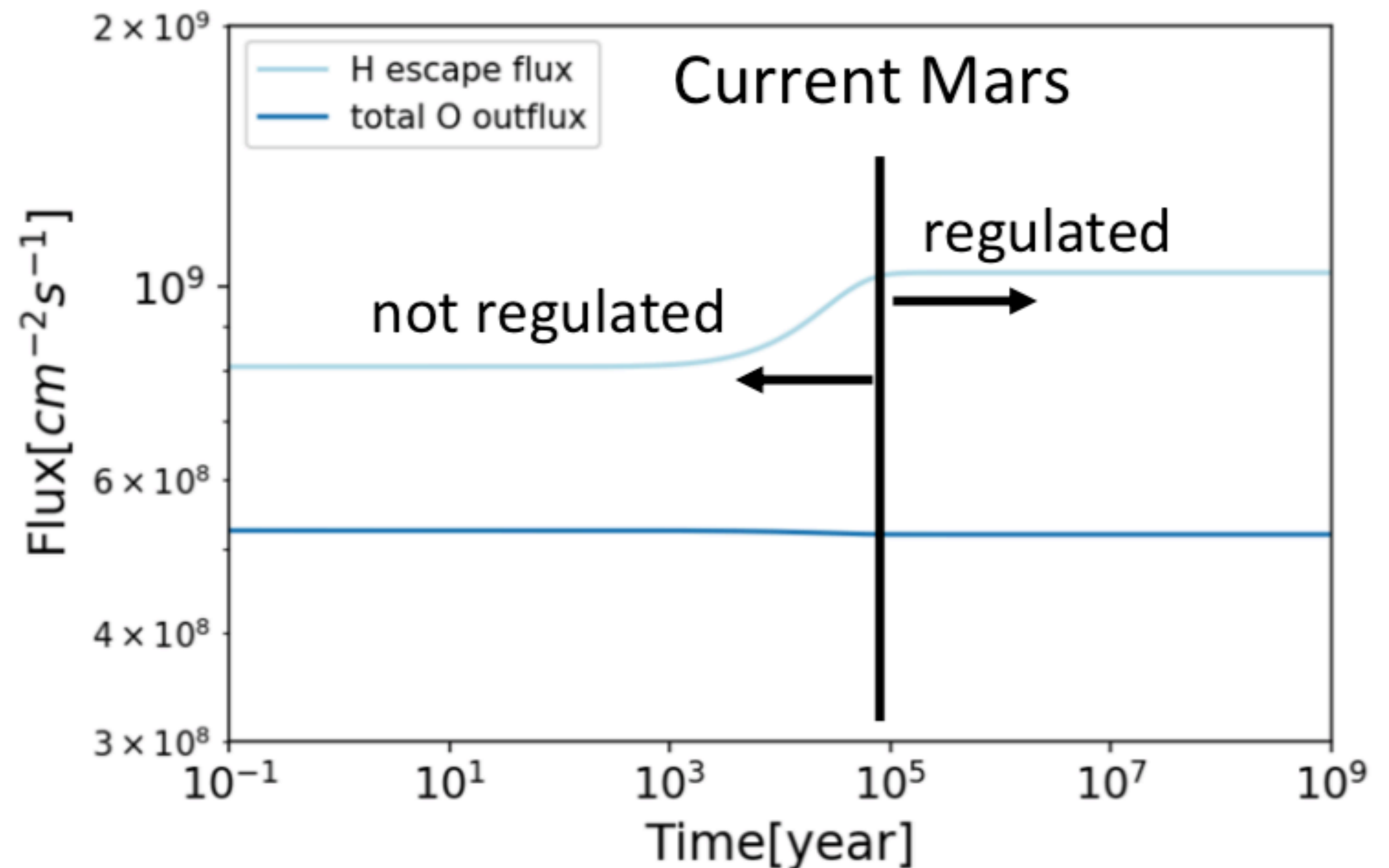
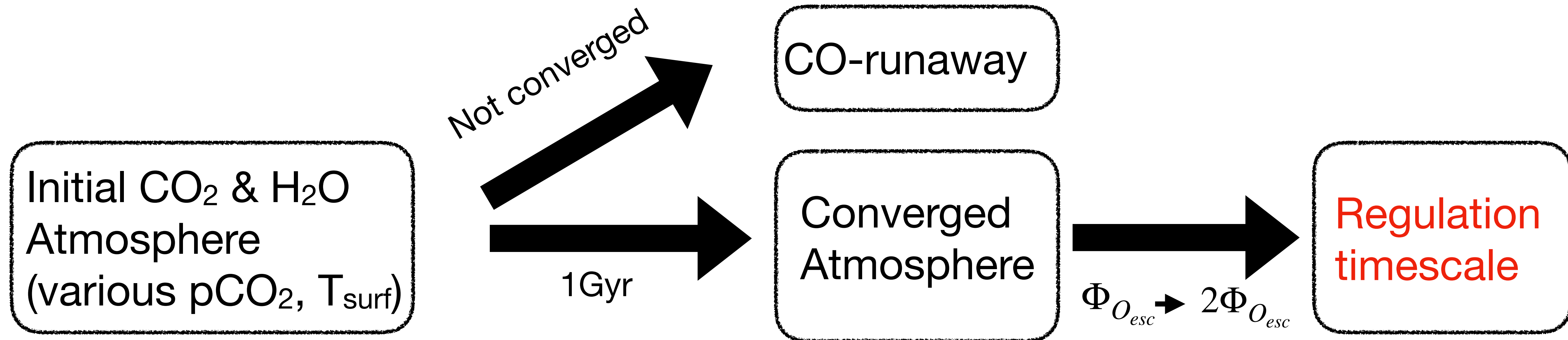
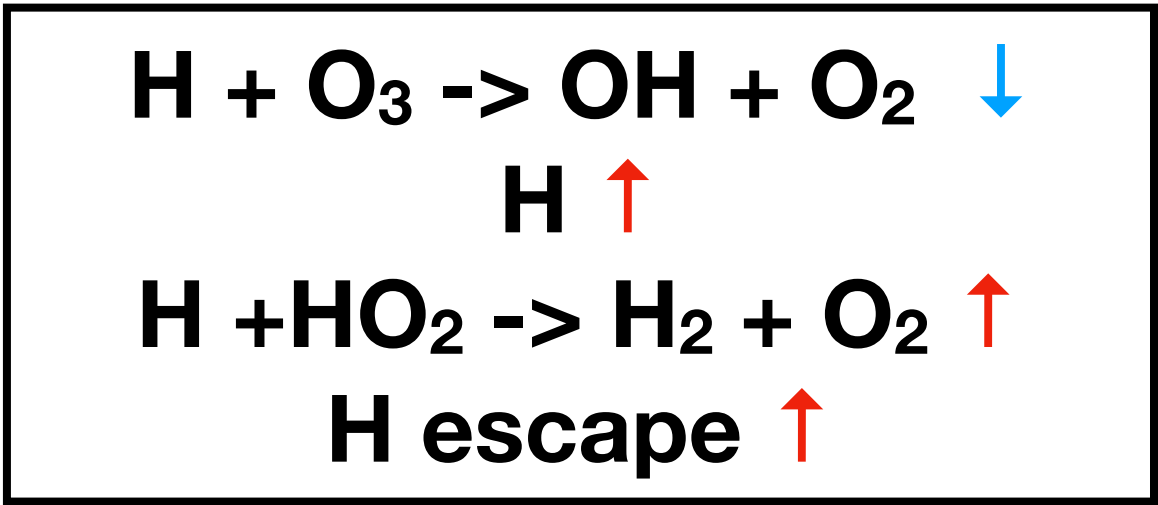
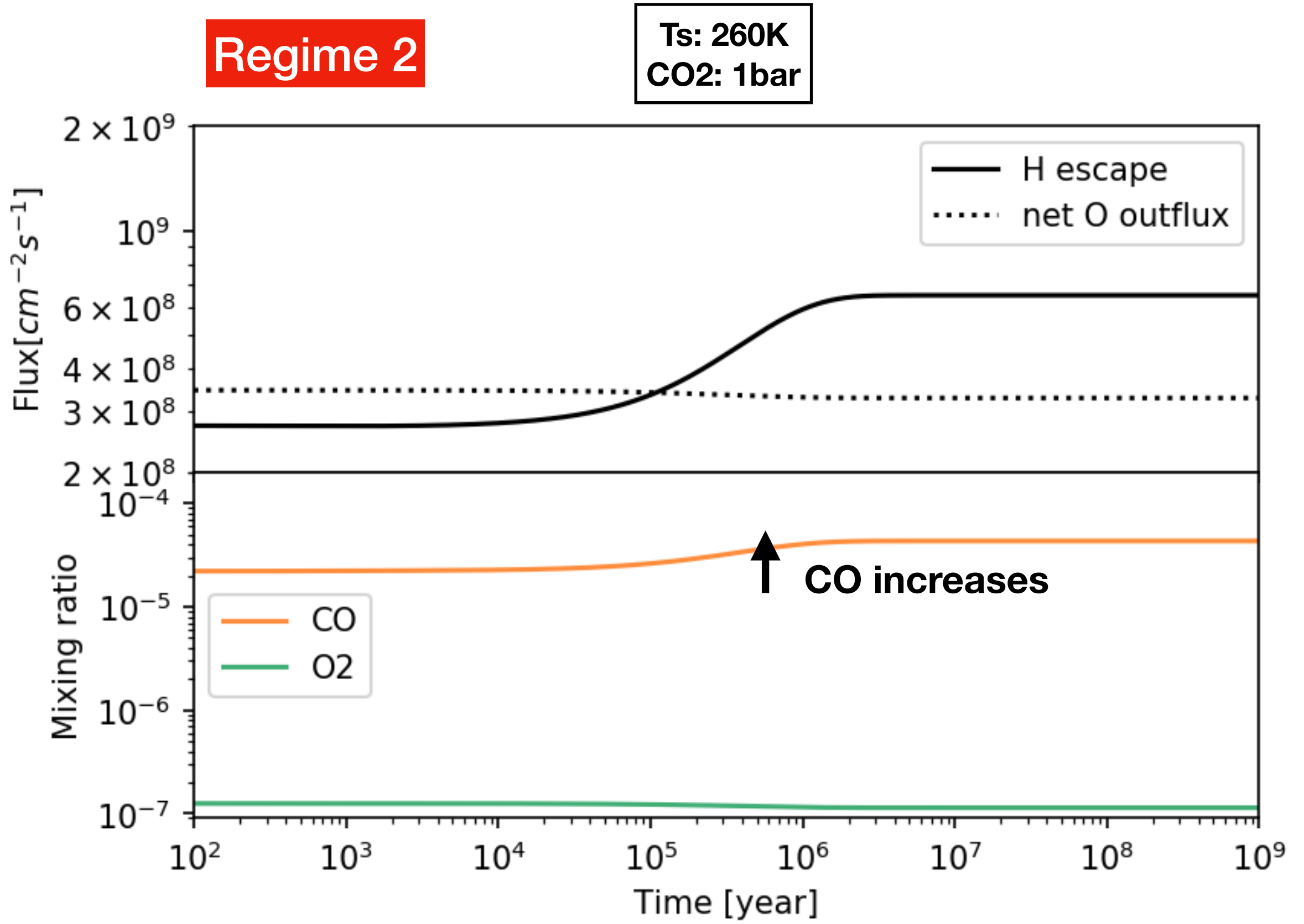
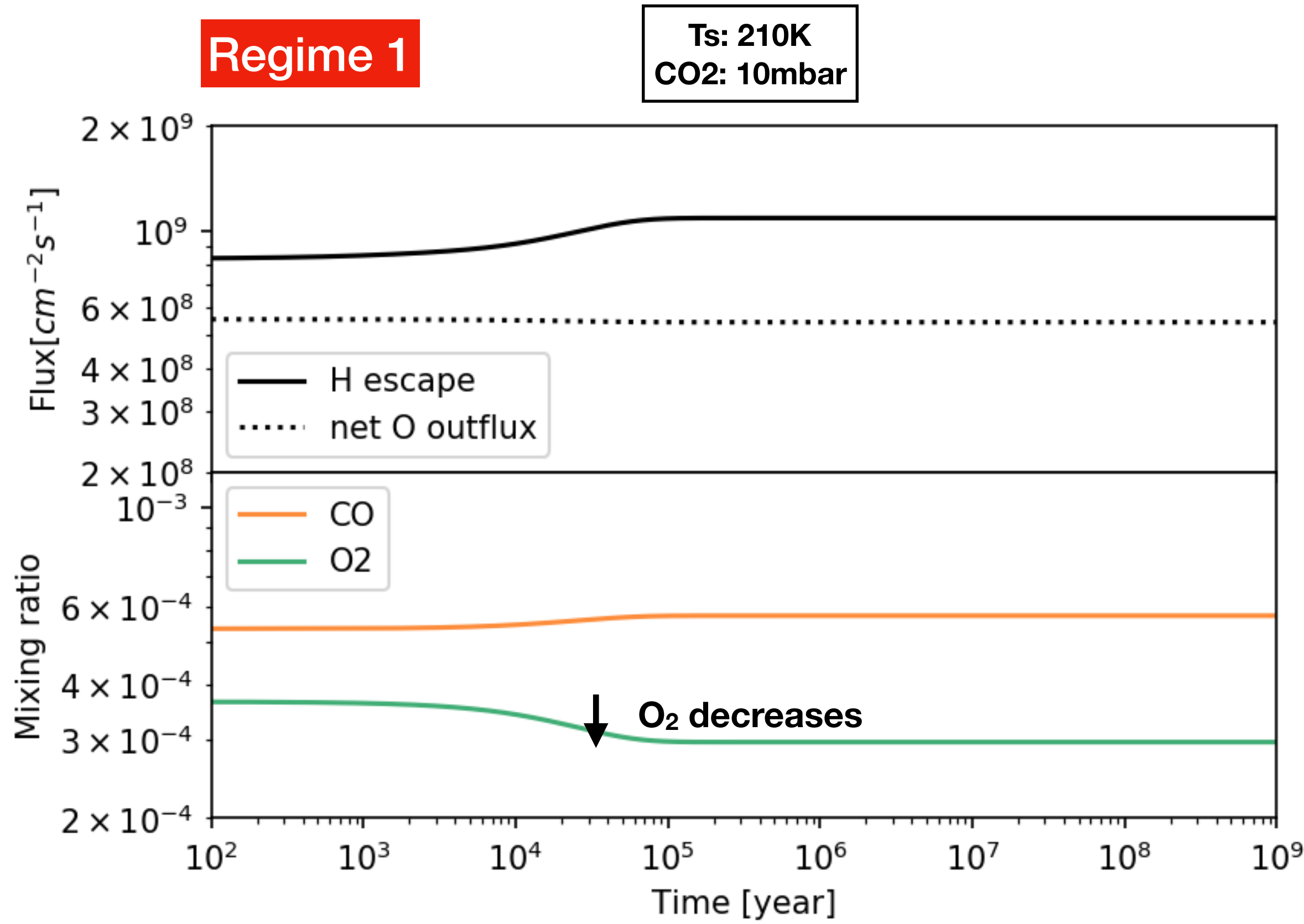


Fig.  $\Phi_{O_{esc}} = 1.2 \times 10^8 \rightarrow 2.4 \times 10^8 \text{ cm}^{-2} \text{s}^{-1}$  at time zero under current Martian atmospheric condition

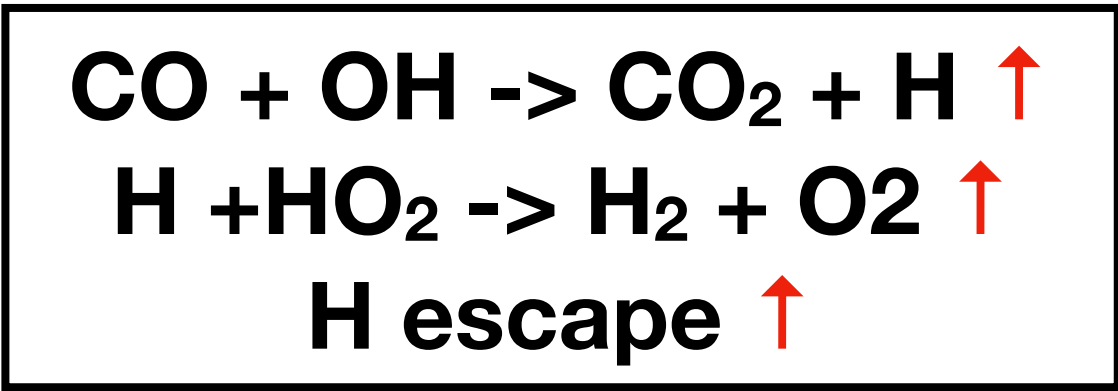
# Summary of what we did



# Two Regulation regimes

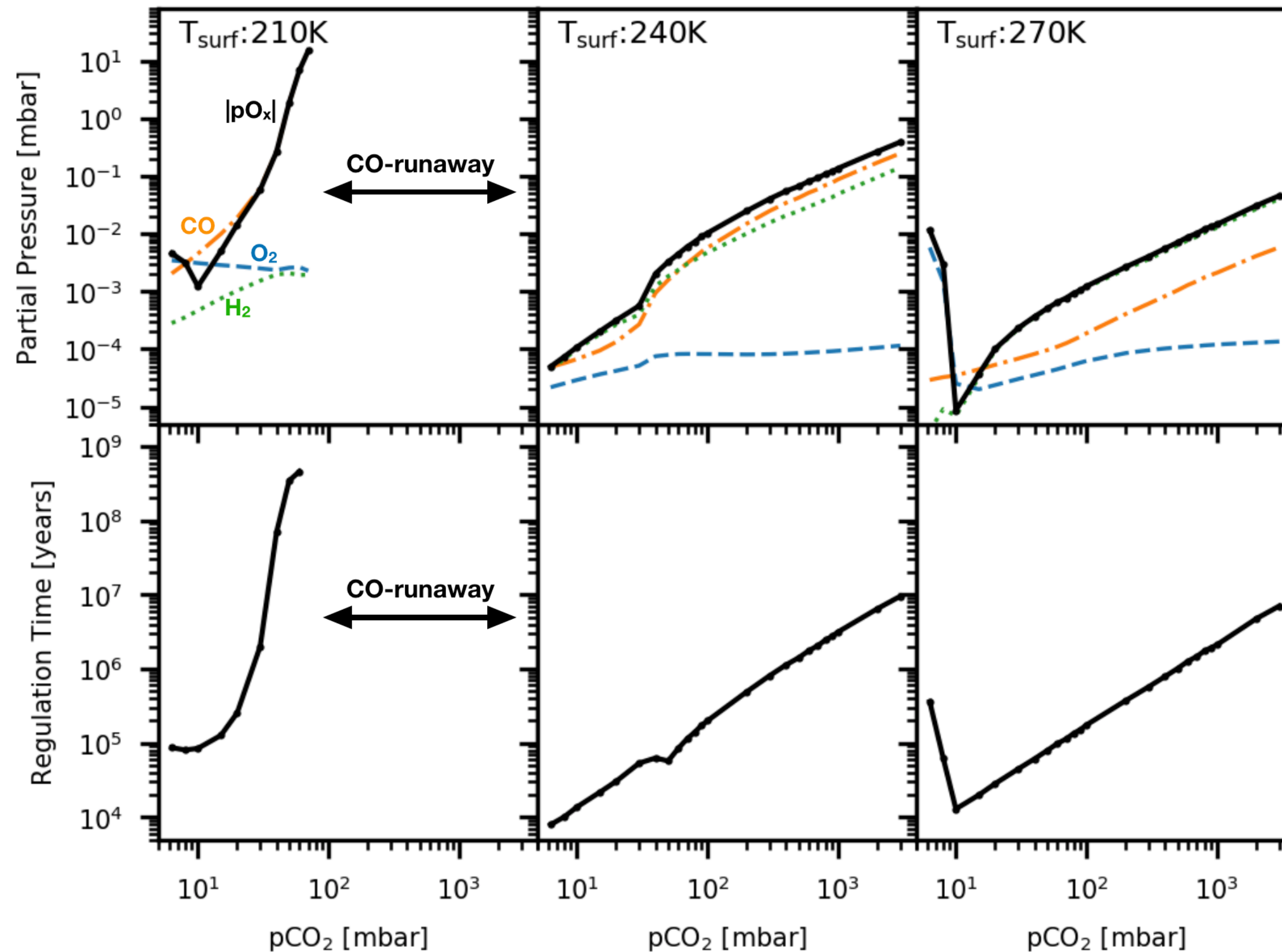


Timescale  
 $\propto \text{total } O_2 / \Phi_{O_{esc}}$



Timescale  
 $\propto \text{total } CO / \Phi_{O_{esc}}$

# Self-regulation timescales



Net oxidation:

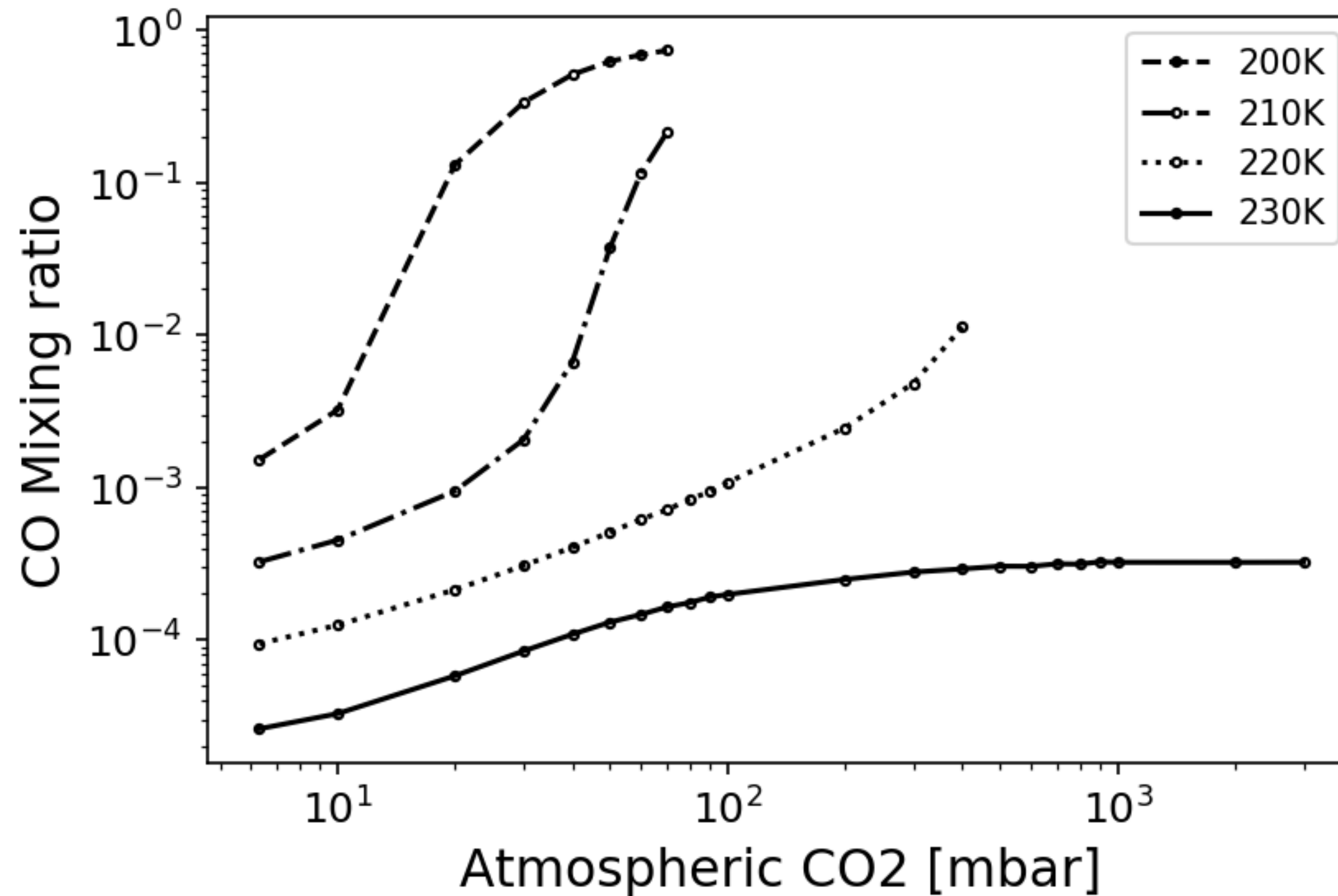
$$pOx = 2pO_2 - pCO - pH_2$$

[Zahnle et al., 2008]

- Self-regulation timescale is likely to be controlled by  $|p\text{O}_x|$ .
- The atmosphere with more  $\text{CO}_2$  on early Mars has the longer timescale than the present.

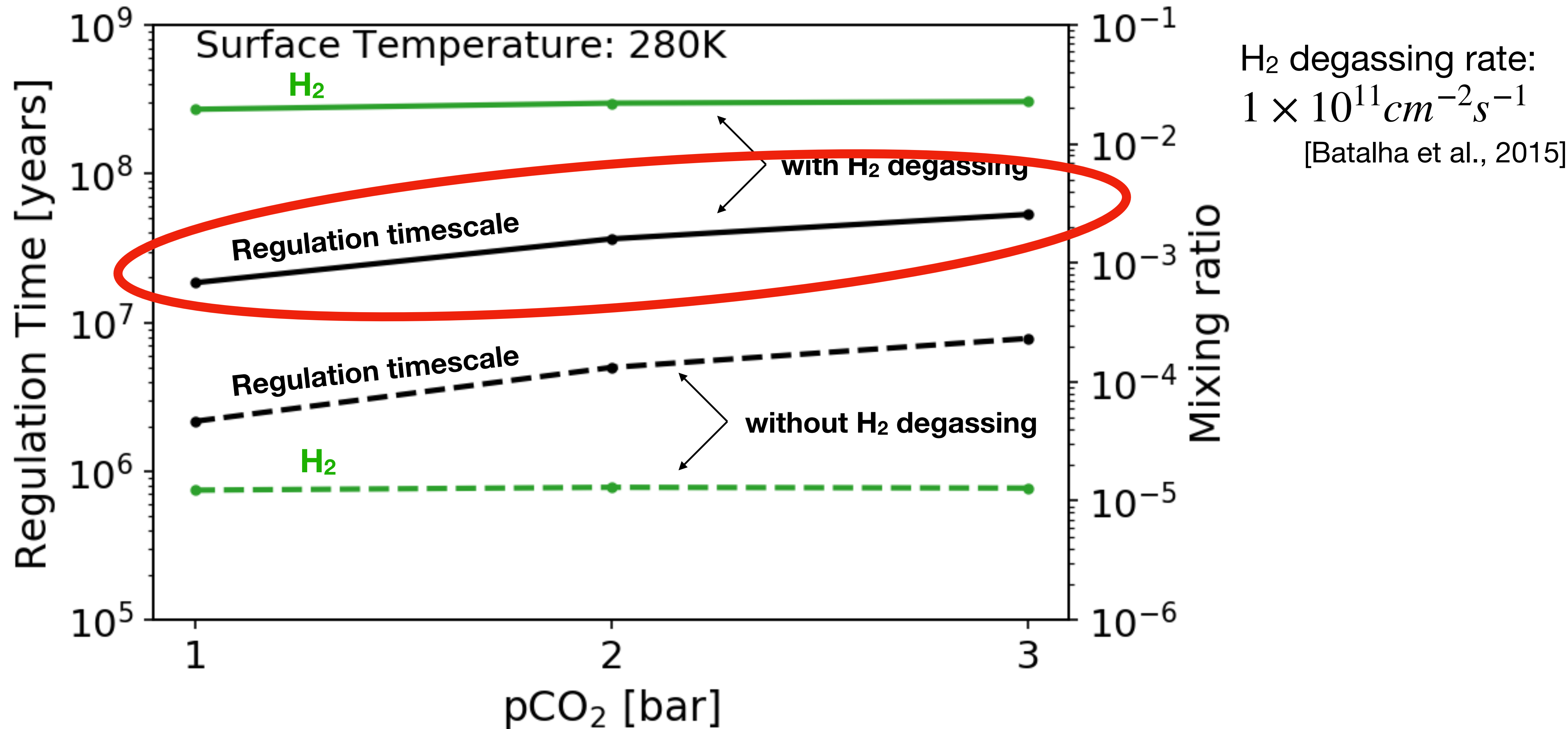


# CO possibility on early Mars



- CO<sub>2</sub> atmosphere of 100mbar is converted to CO-main atmosphere in several hundred million years with  $T_{\text{surf}}=200\text{K}$  in our calculation.
- **CO-main atmosphere of about 50-150mbar might be possible around 3 Ga**

# Implication for H<sub>2</sub>-rich CO<sub>2</sub> atmosphere scenario of early Mars



**The redox state of H<sub>2</sub>-CO<sub>2</sub> atmosphere is less stable than current Mars  
 -> The imbalance of H and O losses may have a large impact on its redox state.**

# Summary

- We calculated self-regulation timescales for various atmospheric conditions.
- The regulation timescale is likely to be controlled by net oxidation(|pOx|)
- Denser CO<sub>2</sub> atmospheres on early Mars are less redox-stable (longer timescale) than present-day Mars
- CO-main atmosphere might be possible around 3 Ga.
- The imbalance of H and O outfluxes ratio in H<sub>2</sub>-CO<sub>2</sub> atmosphere might have a large effect on its H<sub>2</sub> abundance.